A satellite image of a tropical cyclone, showing a distinct eye and spiral cloud bands over a dark blue ocean. The text is overlaid on the top half of the image.

Investigating Tropical Cyclone Intensity Change using WRF simulations, In-Situ Field Data and Satellite and Aircraft Remote Sensing Observations

**G. M. McFarquhar, B. F. Jewett, S. W. Nesbitt, M. Gilmore, D. Harnos,
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Acknowledgments

S. Braun

A. Heymsfield

G. Heymsfield

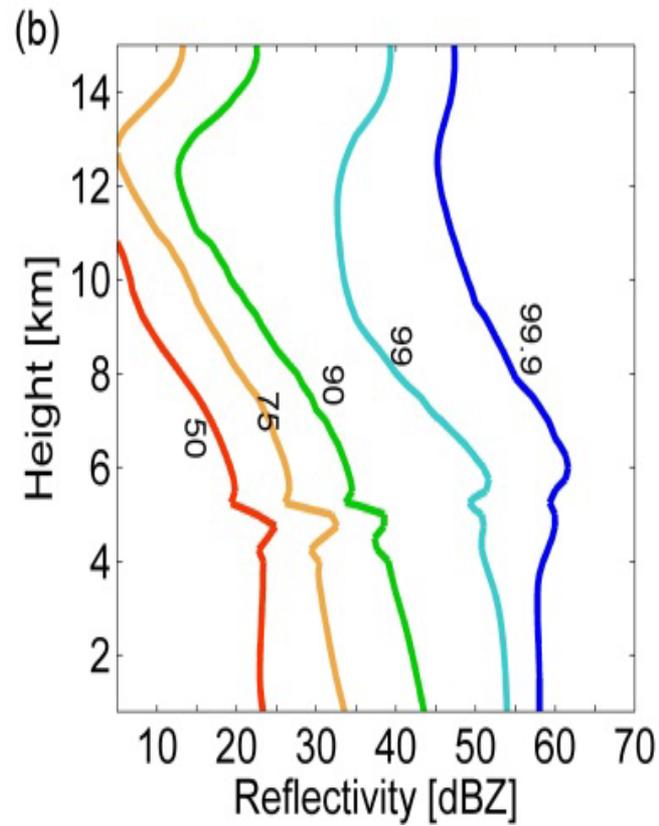
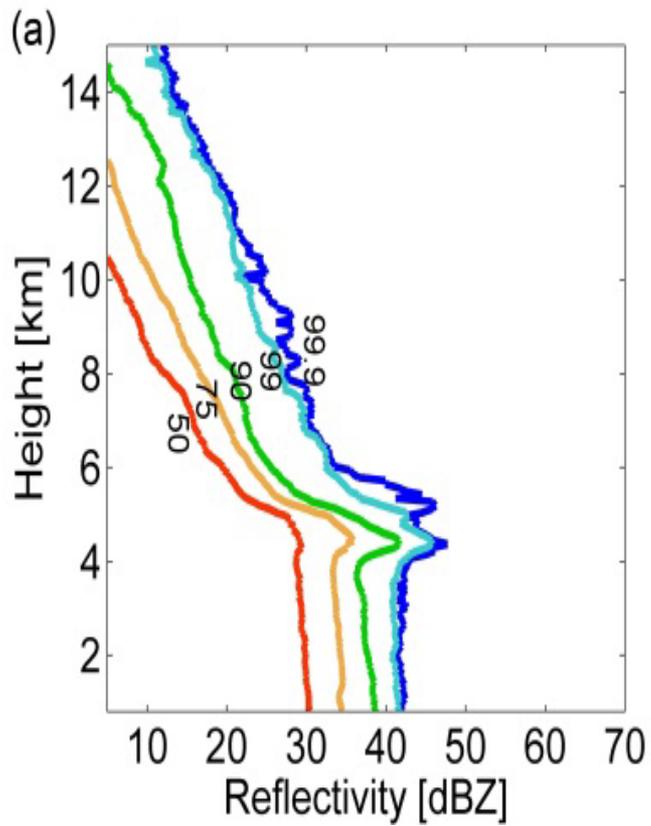
J. Molinari

R. Rogers

R. Kakar

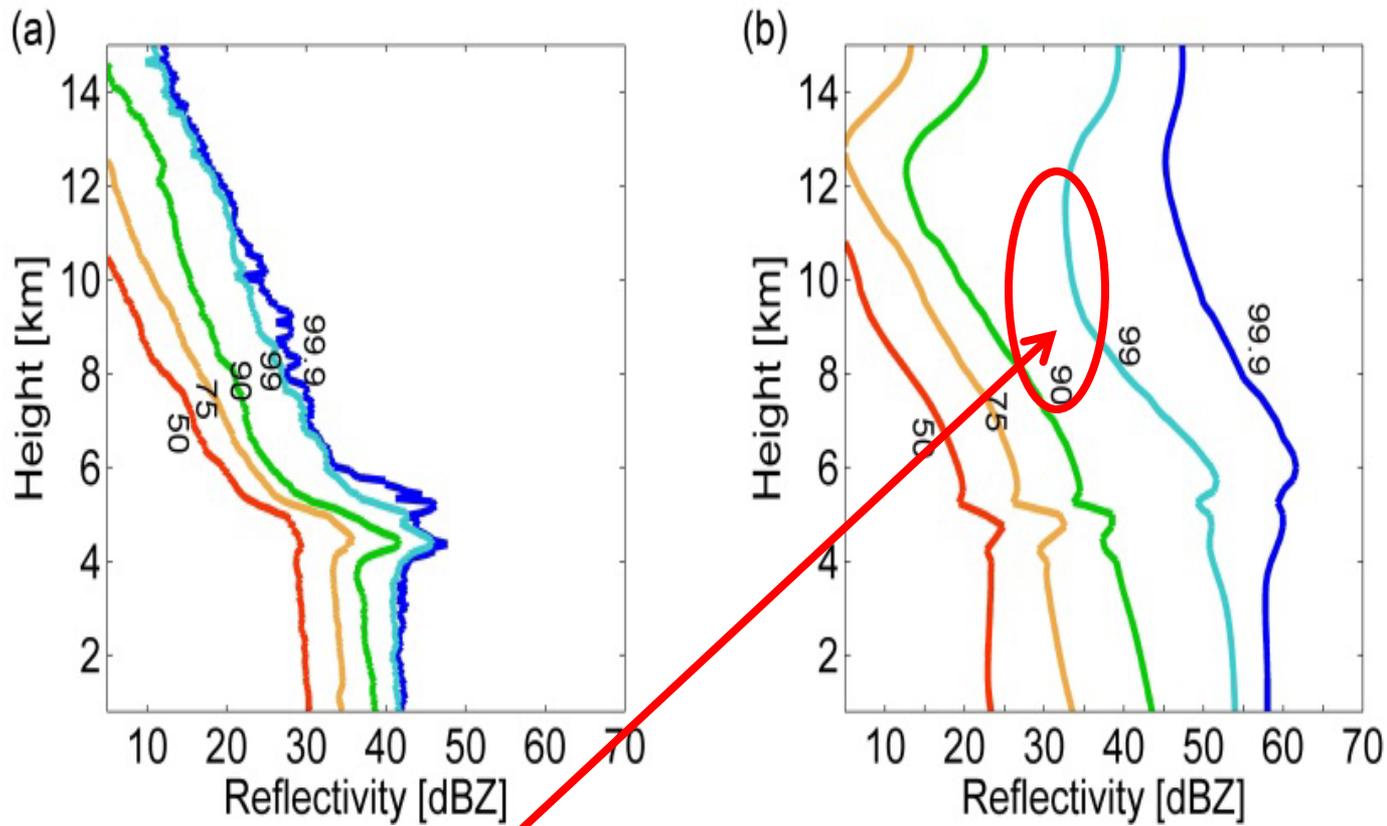
Research Topics

- 1. Examine mechanisms driving intensity change of Dennis (2005), Earl (2010) and Ike (2008) using WRF simulations & observations.**
- 2. Assess how dust in SAL acting as CCN impacts TC intensity using ensemble of simulations.**
- 3. Using in-situ data to improve model representations of hydrometeor SDs as gamma functions**
- 4. Construct comprehensive set of space borne microwave & radar data to examine role of varying TC morphology on RI.**

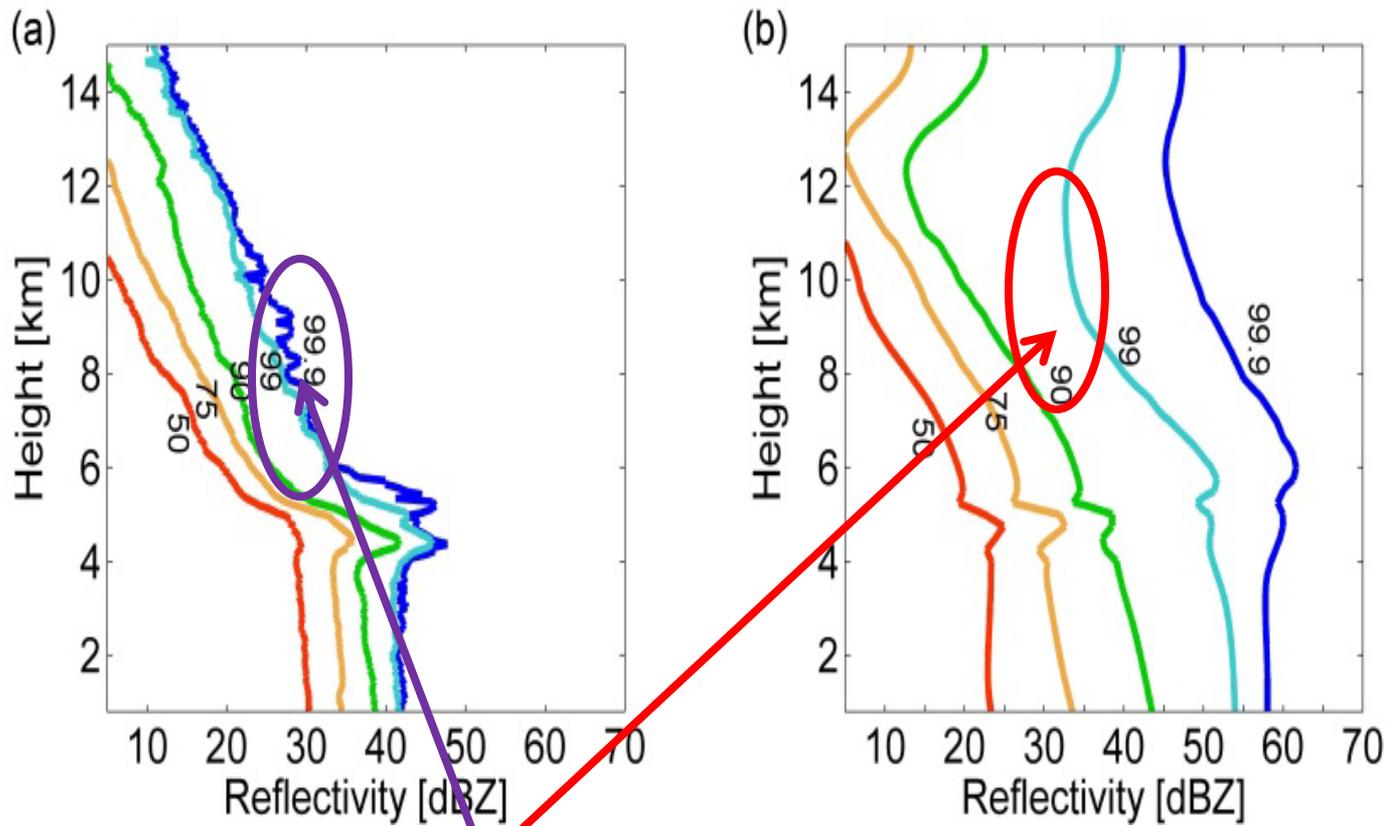


WRF simulation

- 1-km grid spacing, 55 levels
- 1-km data saved @ 2 minutes

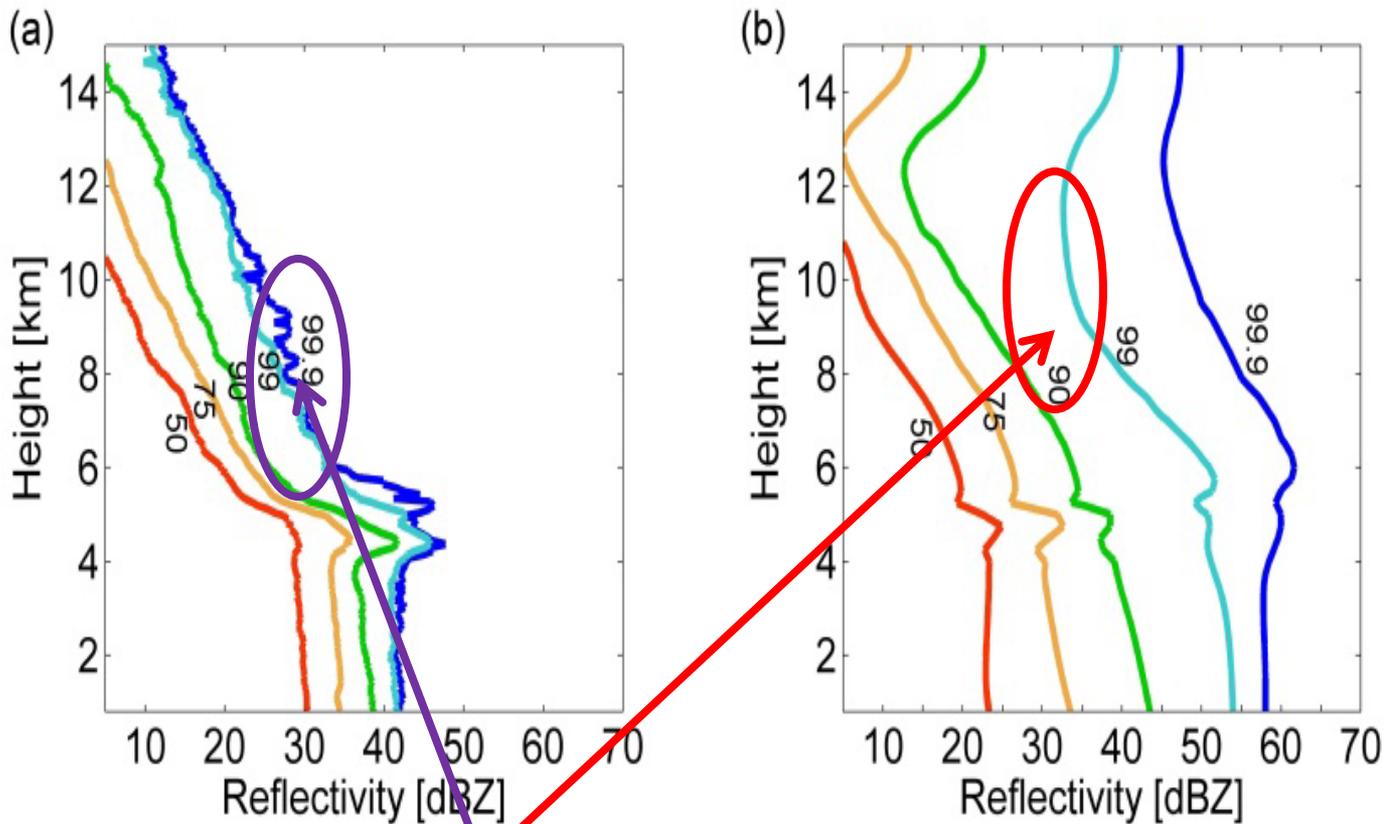


3.0% of Z exceed 25 dBZ at 10 km in simulation



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< 0.1 % of Z > 25 dBZ in EDOP data



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< 0.1 % of Z > 25 dBZ in EDOP data

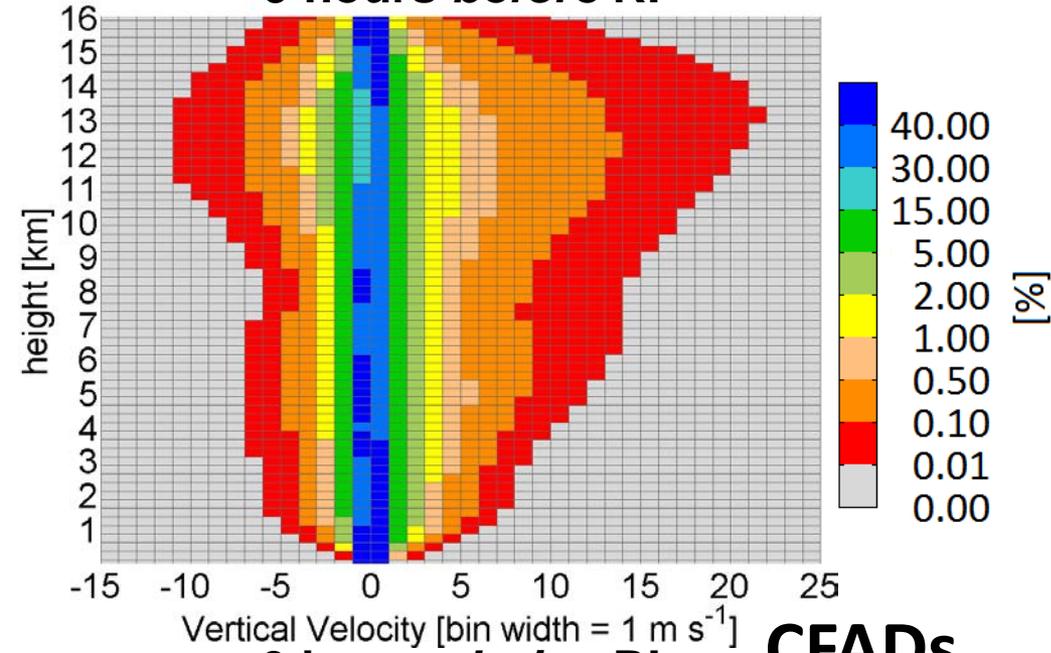
Frequency of largest Z above melting layer over-predicted



WRF Simulations of Hurricane Dennis

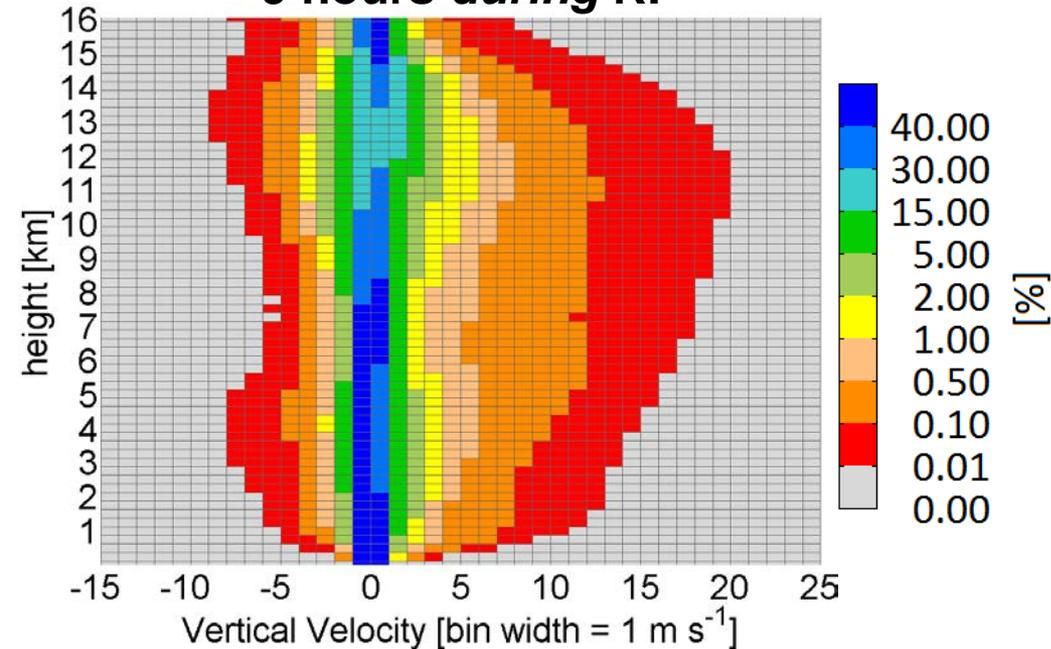
**Cumulative
CFADs within
75 km before &
during RI**

9 hours *before* RI



CFADs

9 hours *during* RI

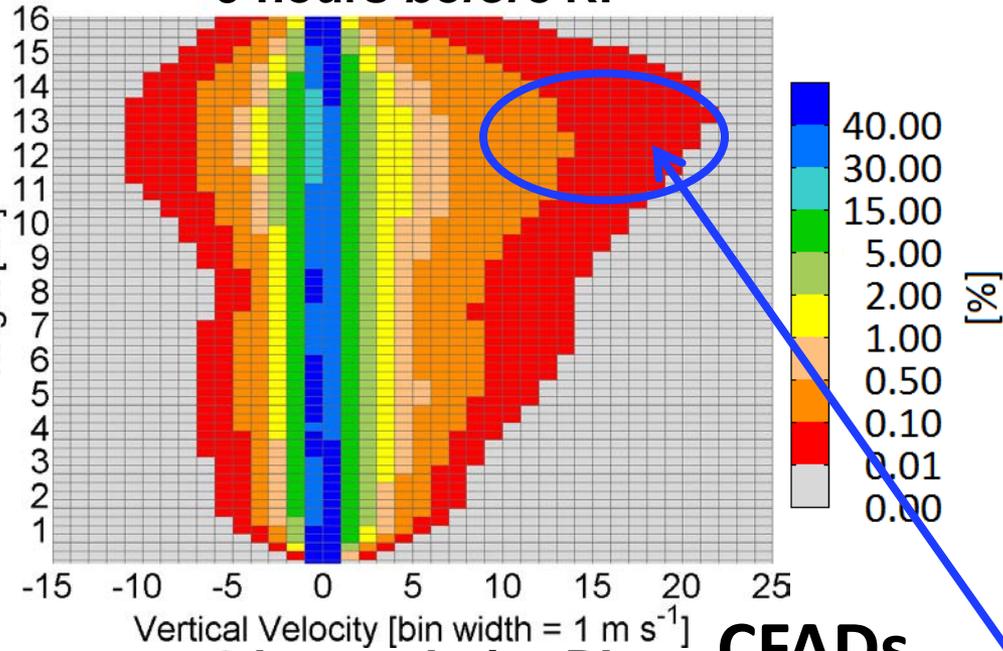


WRF Simulations of Hurricane Dennis

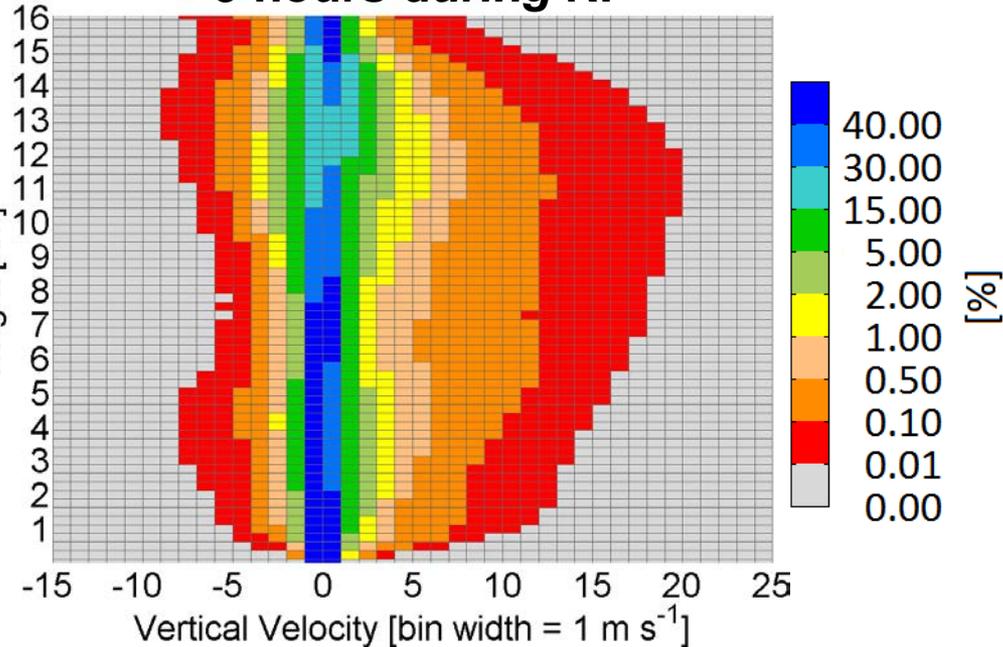
Cumulative CFADs within 75 km before & during RI

Distribution broader before RI for $z > 11$ km,

9 hours before RI



9 hours during RI



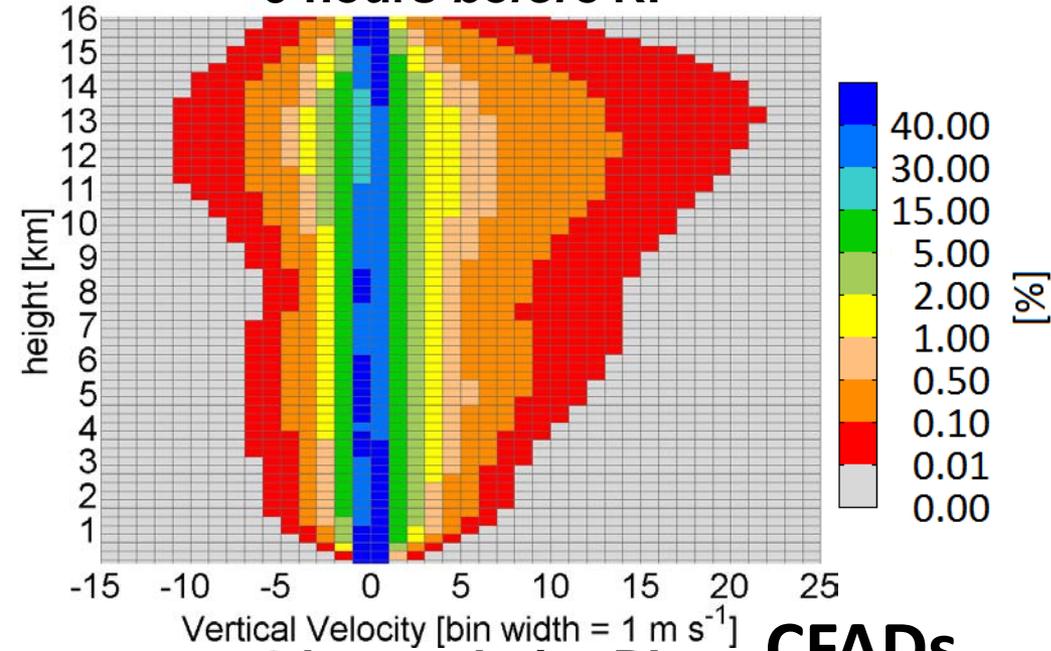
CFADs

WRF Simulations of Hurricane Dennis

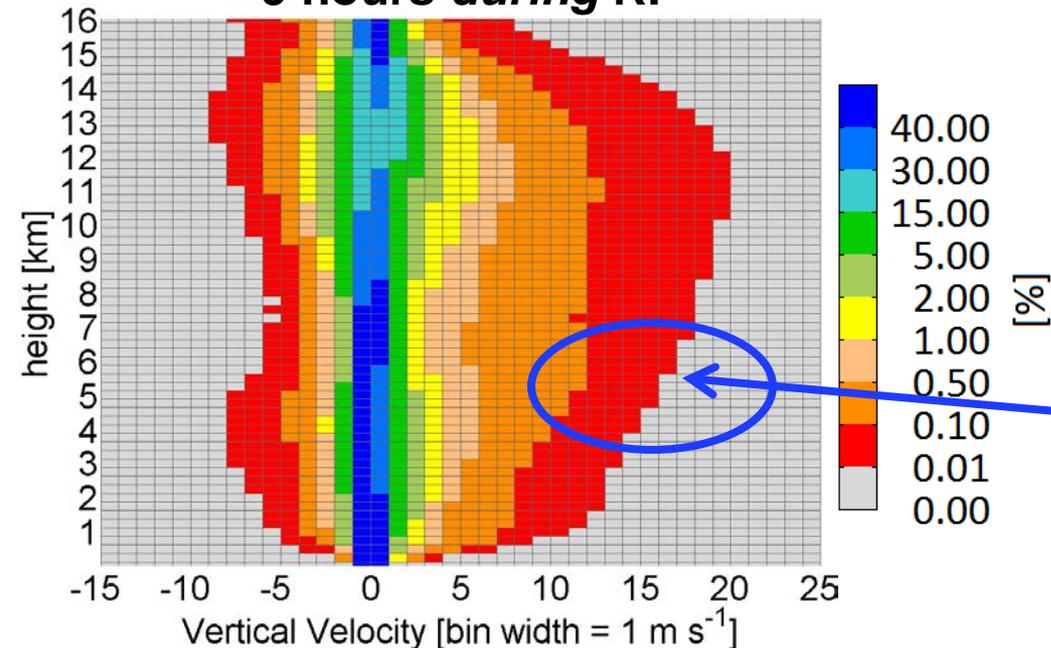
Cumulative CFADs within 75 km before & during RI

Distribution broader before RI for $z > 11$ km, but broader after RI for $z < 6$ km

9 hours *before* RI



9 hours *during* RI

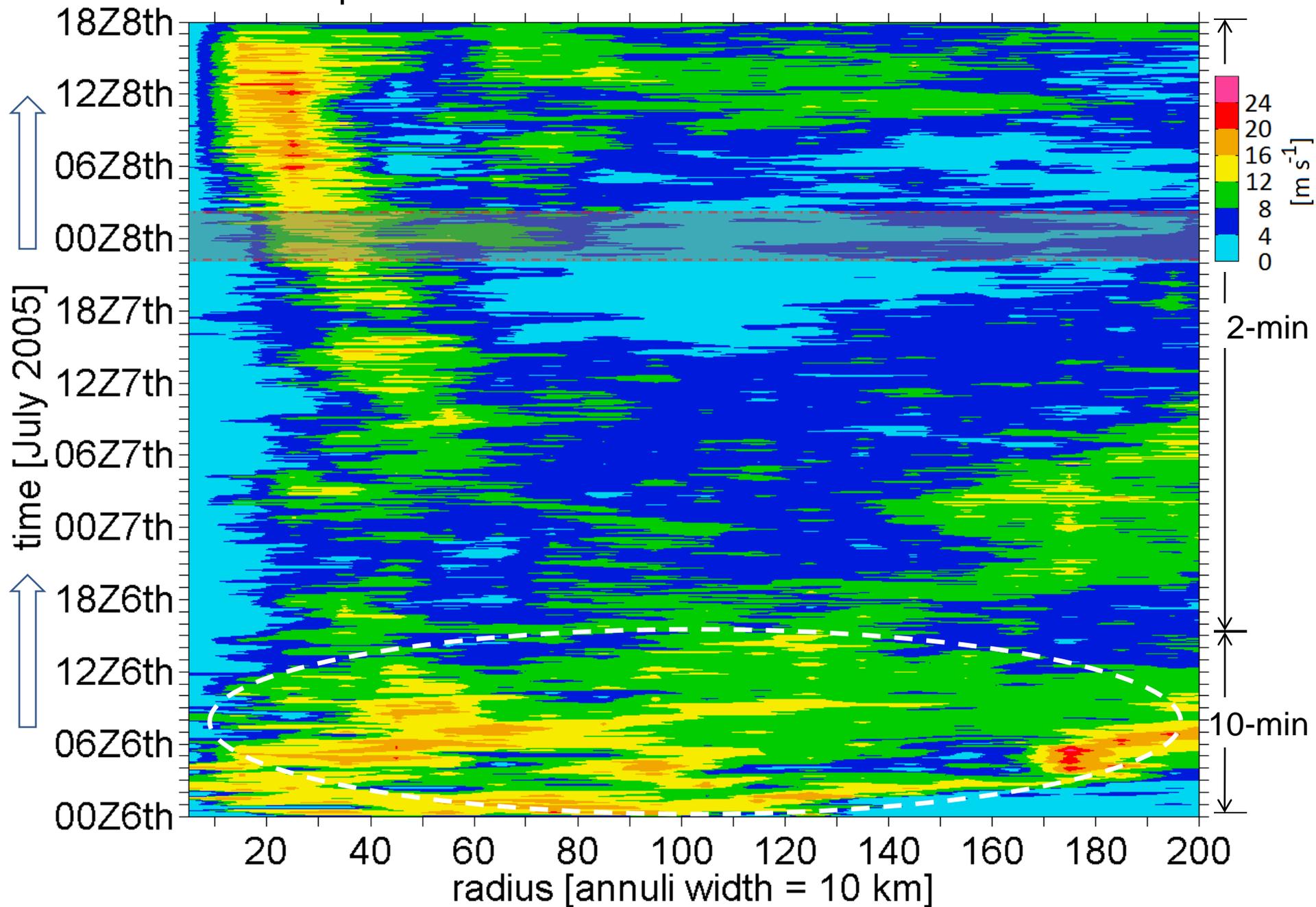


CFADs



99th percentile **VERTICAL VELOCITY**

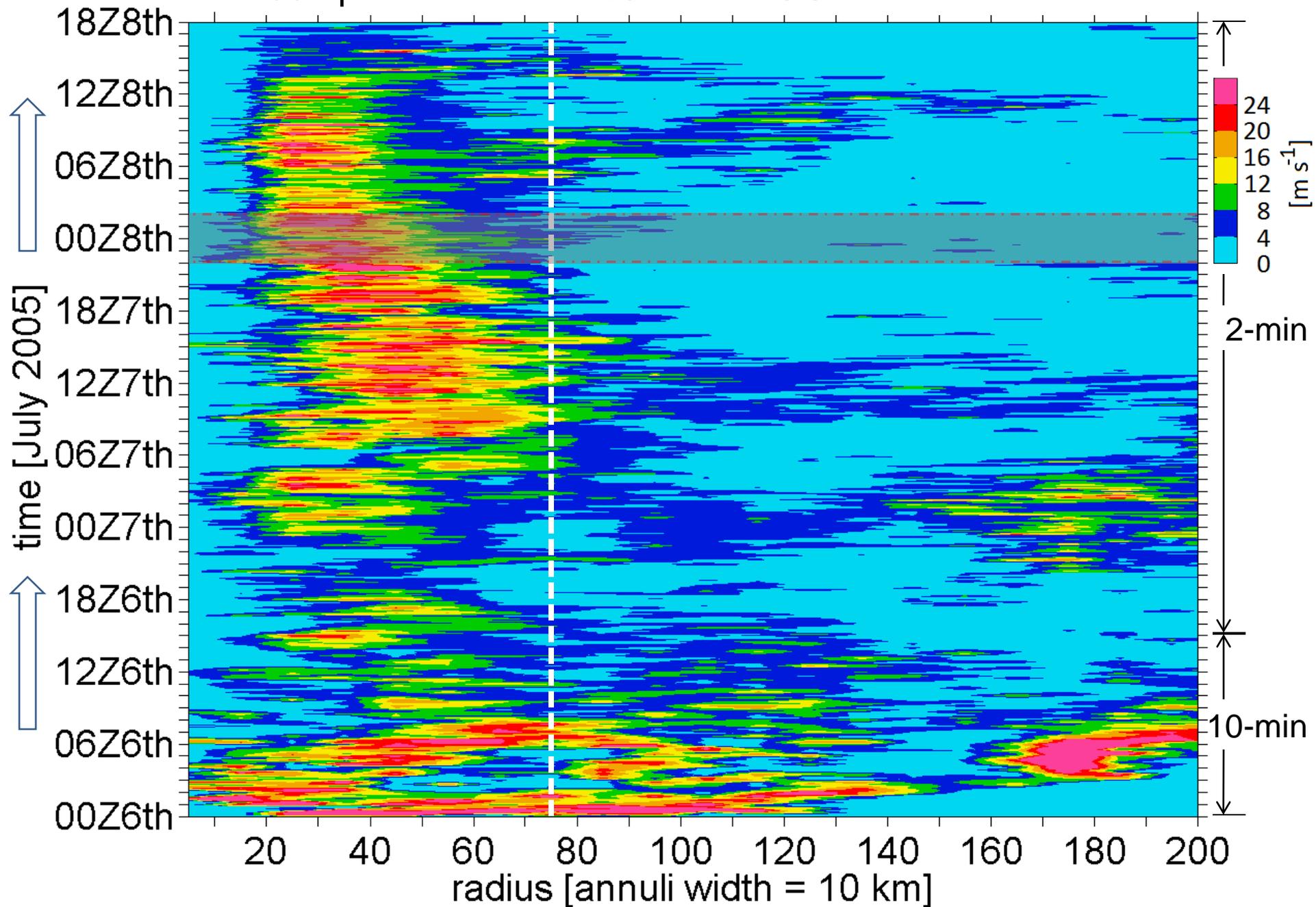
Z=6 km





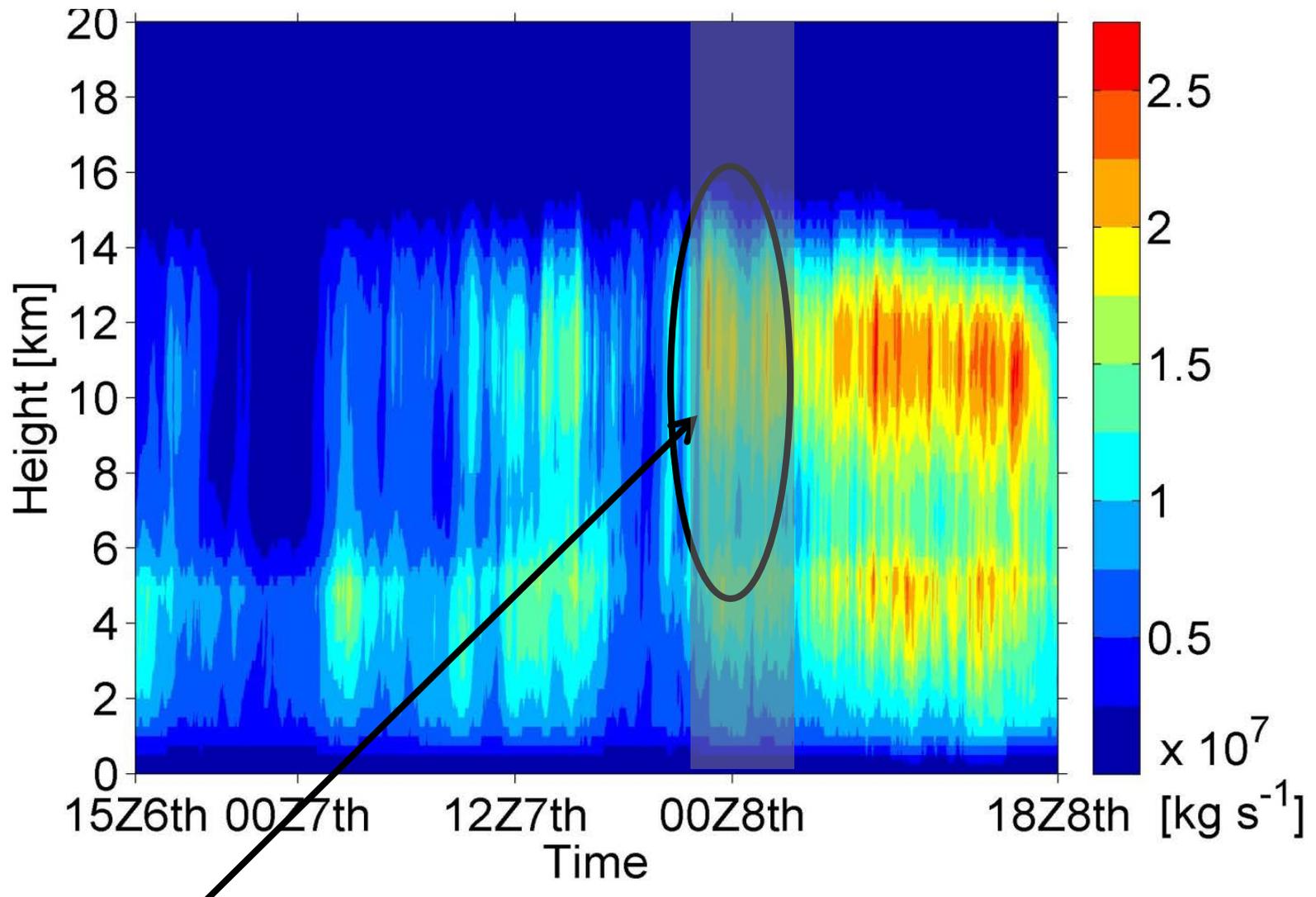
99th percentile **VERTICAL VELOCITY**

Z=14 km



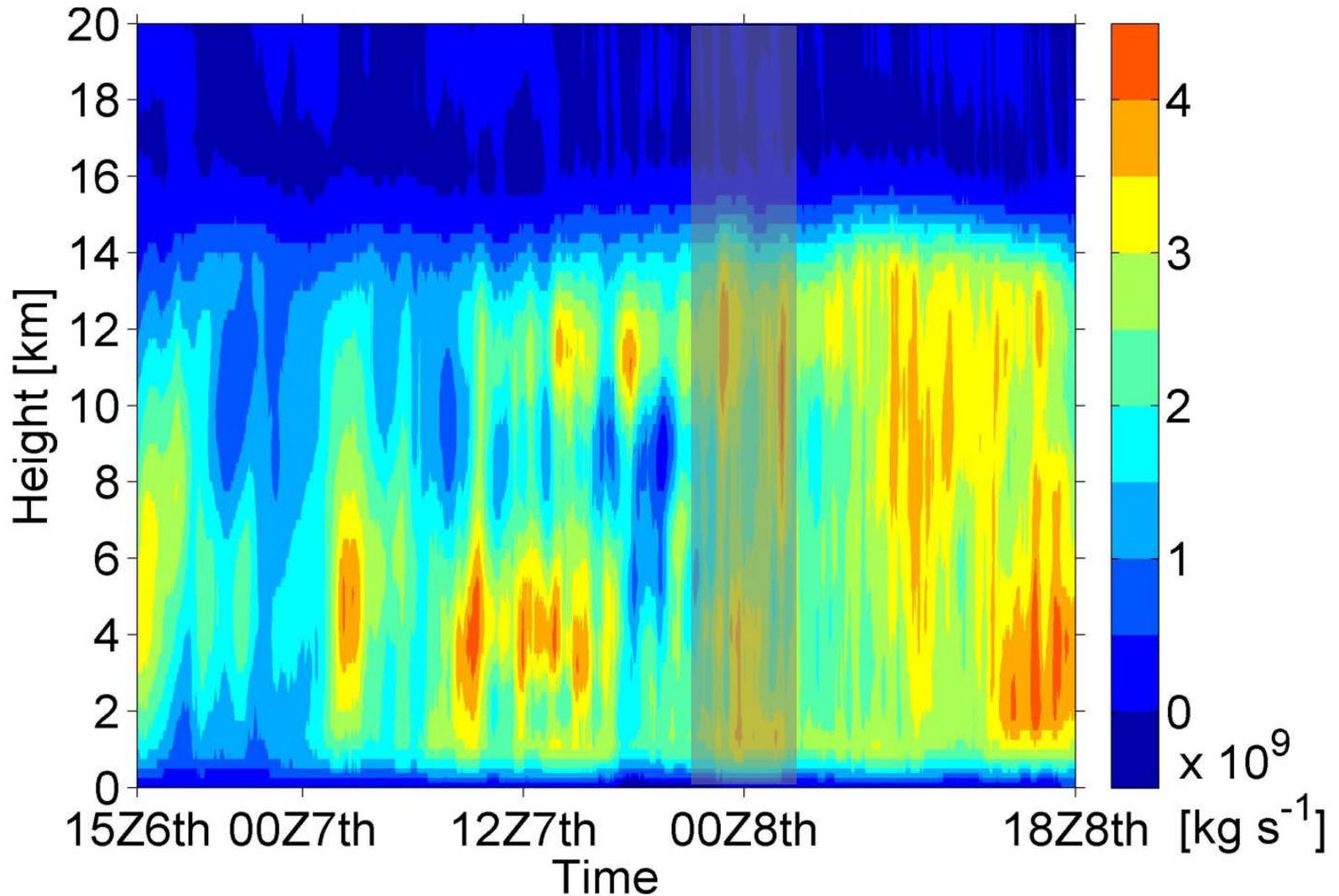
Convective Updraft hydrometeor mass flux

$r < 75 \text{ km}$



**Increase in upward hydrometeor mass flux
at upper levels before RI**

Net air mass flux $r < 75 \text{ km}$

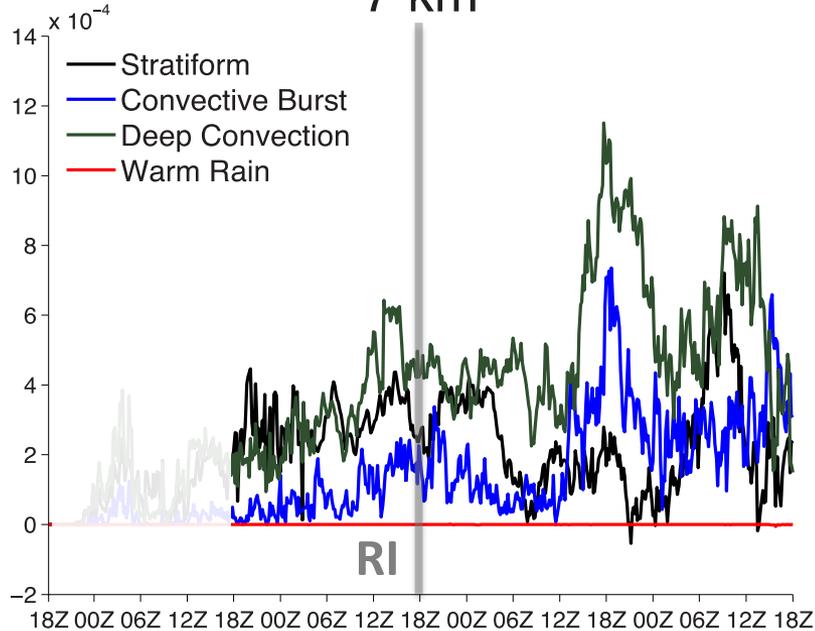


**Increase in air mass flux before RI,
but also occurred well before RI and at lower levels**

WRF: Innermost 1° Diabatic Heating

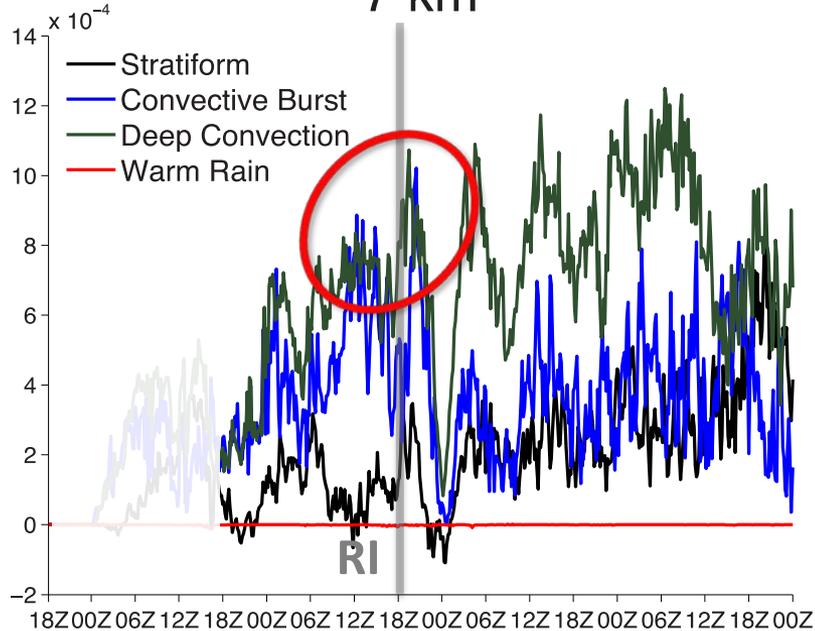
Ike
(Low Shear)

7 km



Earl
(High Shear)

7 km



In high shear, convective bursts cover an order of magnitude less area than deep convection, yet provide more net diabatic heating leading up to RI.

Burst: $W_{mean} \geq 5 \text{ m s}^{-1}$ from 700 - 300 hPa; Rogers (2010)

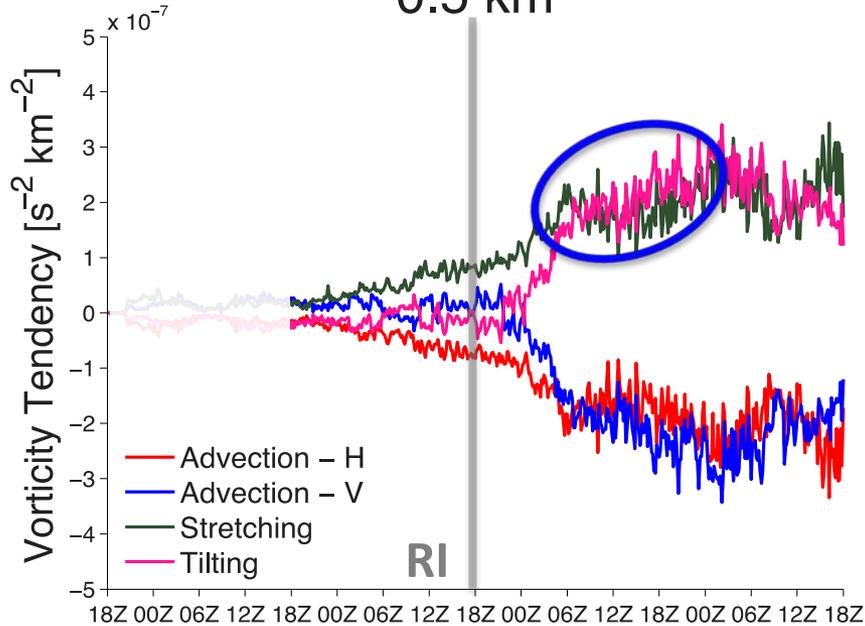
Precipitation Type: Steiner et al. (1995)

WRF: Innermost 1° Vorticity Tendency

Ike

(Low Shear)

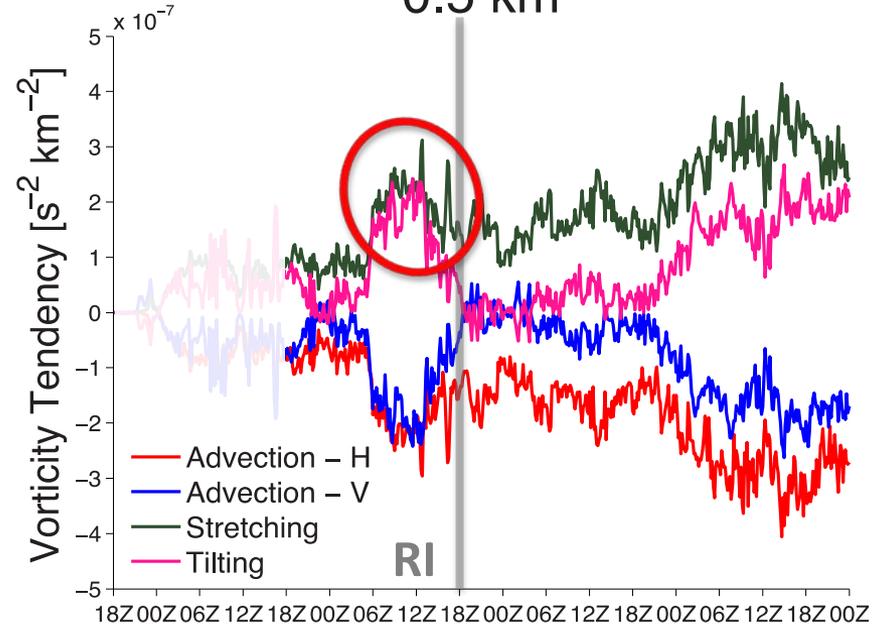
0.5 km



Earl

(High Shear)

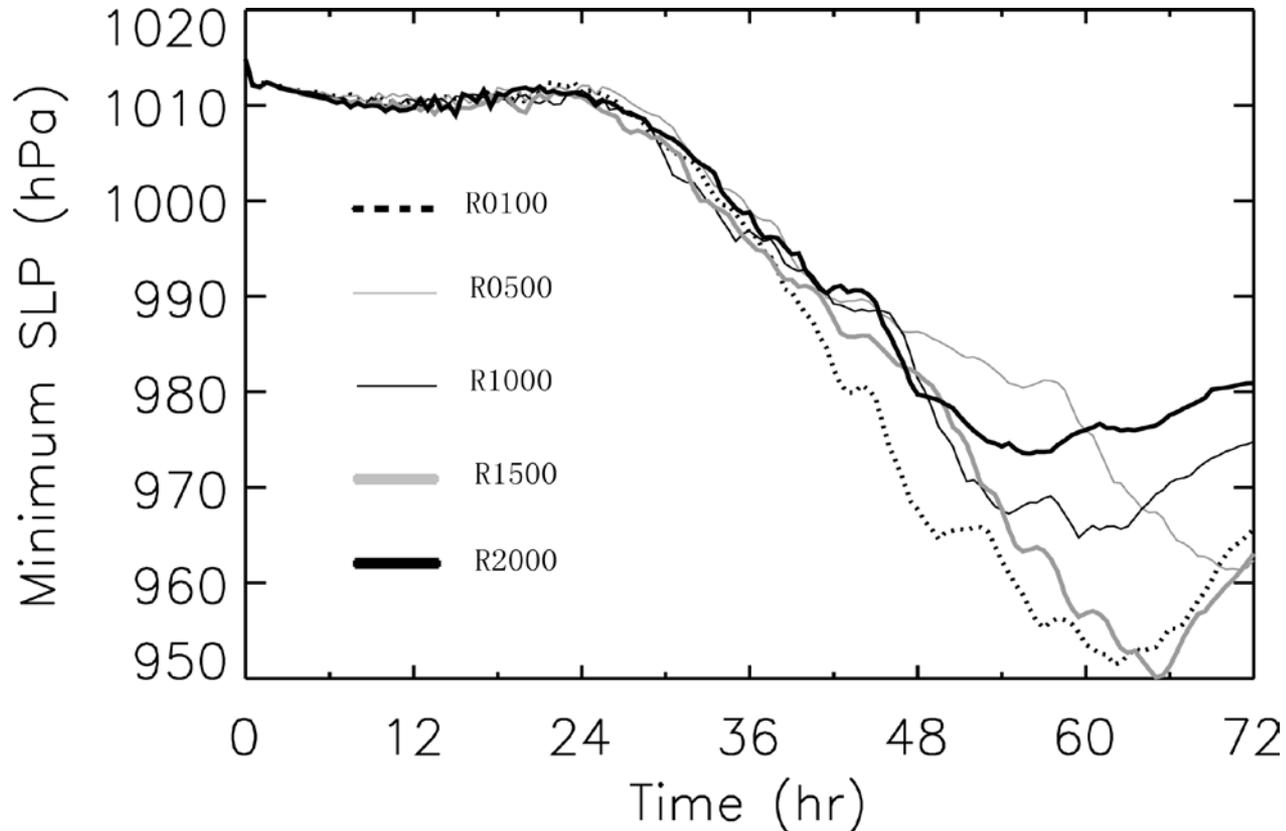
0.5 km



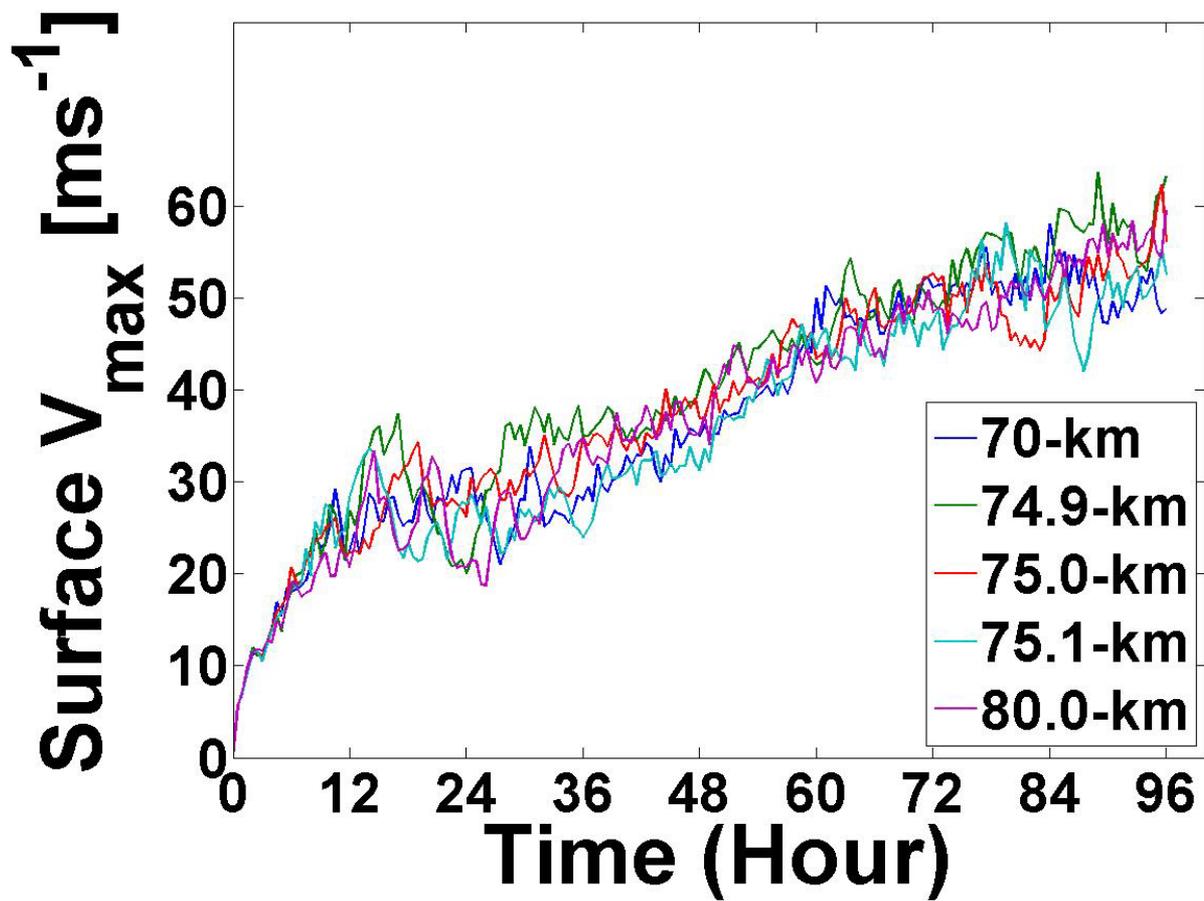
In low shear at low-levels, stretching is the primary contributor leading up to RI, with tilting also contributing later in the RI cycle.

In high shear at low-levels, stretching and tilting contribute nearly the same amount preceding RI, with stretching eventually dominating.

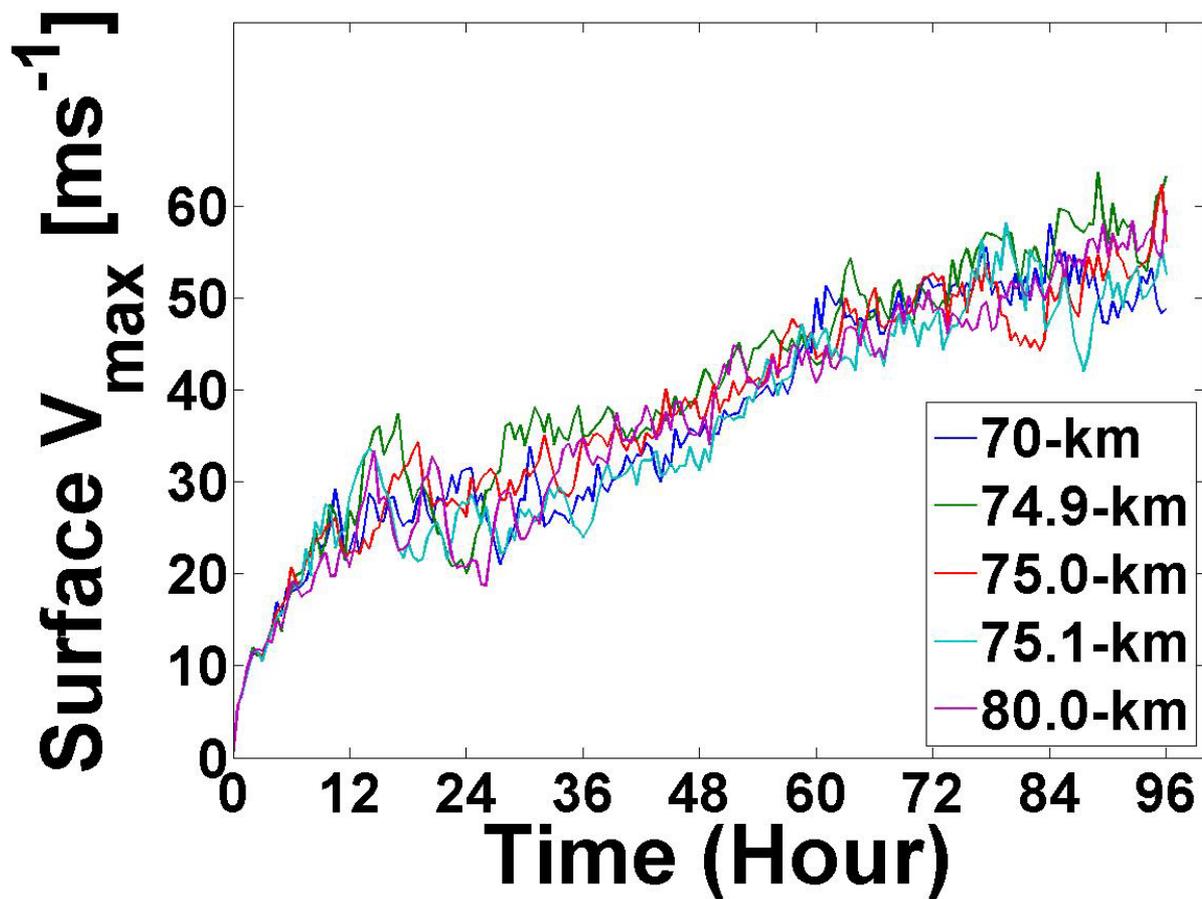
SAL Aerosol Influence on TCs



- **Zhang et al. (2007, 2009) suggested TCs sensitive to variations in CCN number**
 - **TC intensity exhibited non-monotonic response to increases in CCN concentration**

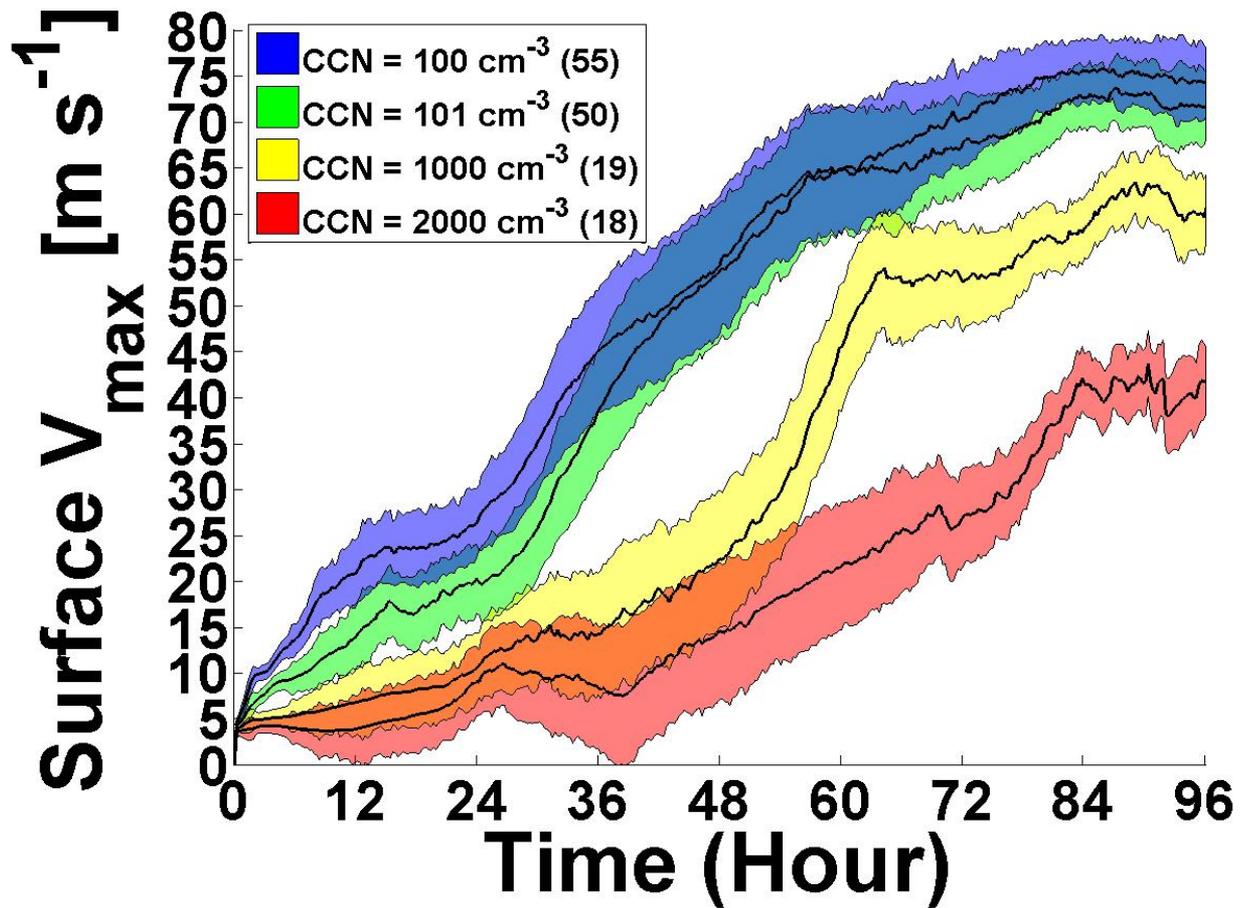


TCs exhibit sensitivity to other ICs, such as RMW of initial vortex

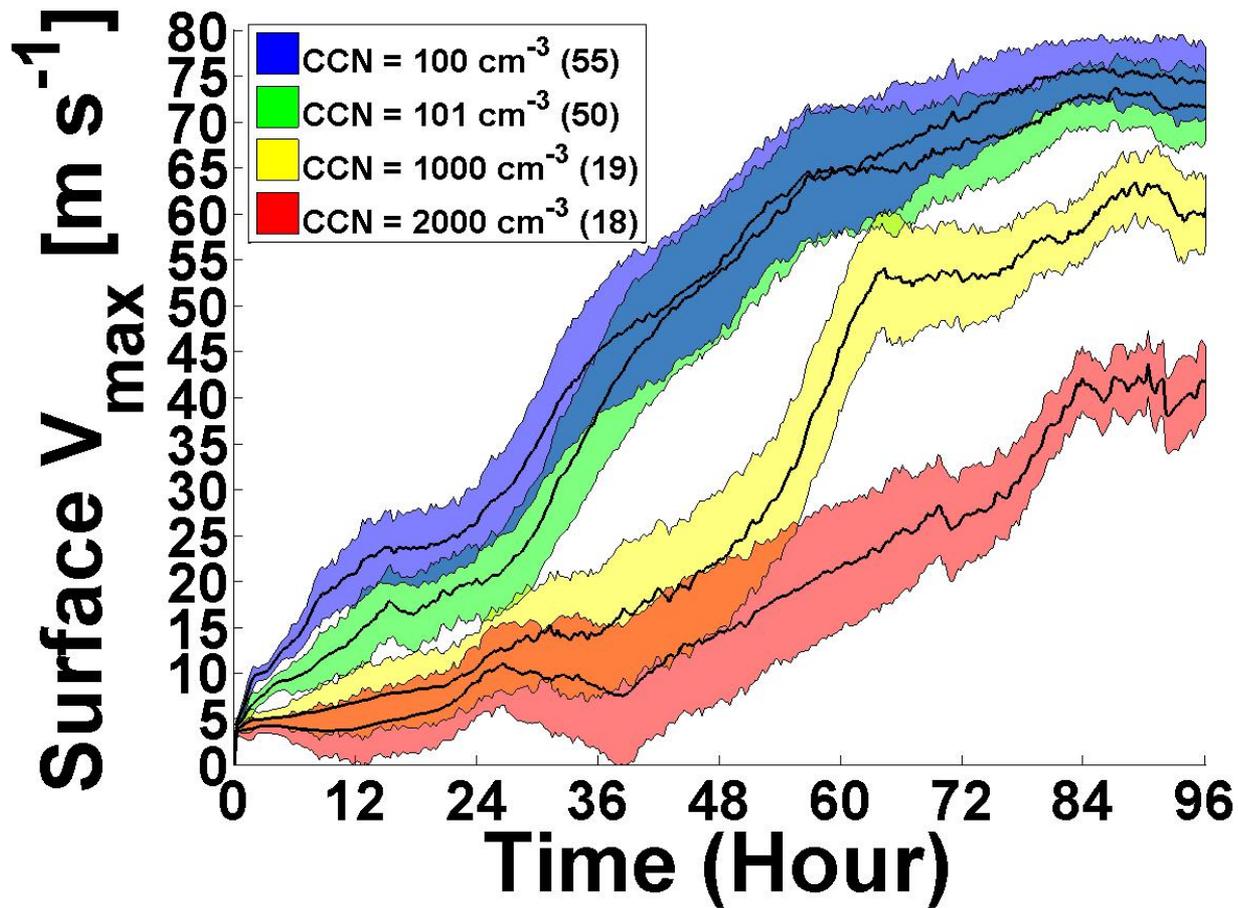


TCs exhibit sensitivity to other ICs, such as RMW of initial vortex

Is non-monotonic response of TCs to CCN caused by *non-linear amplification of noise* associated with perturbations in ICs?

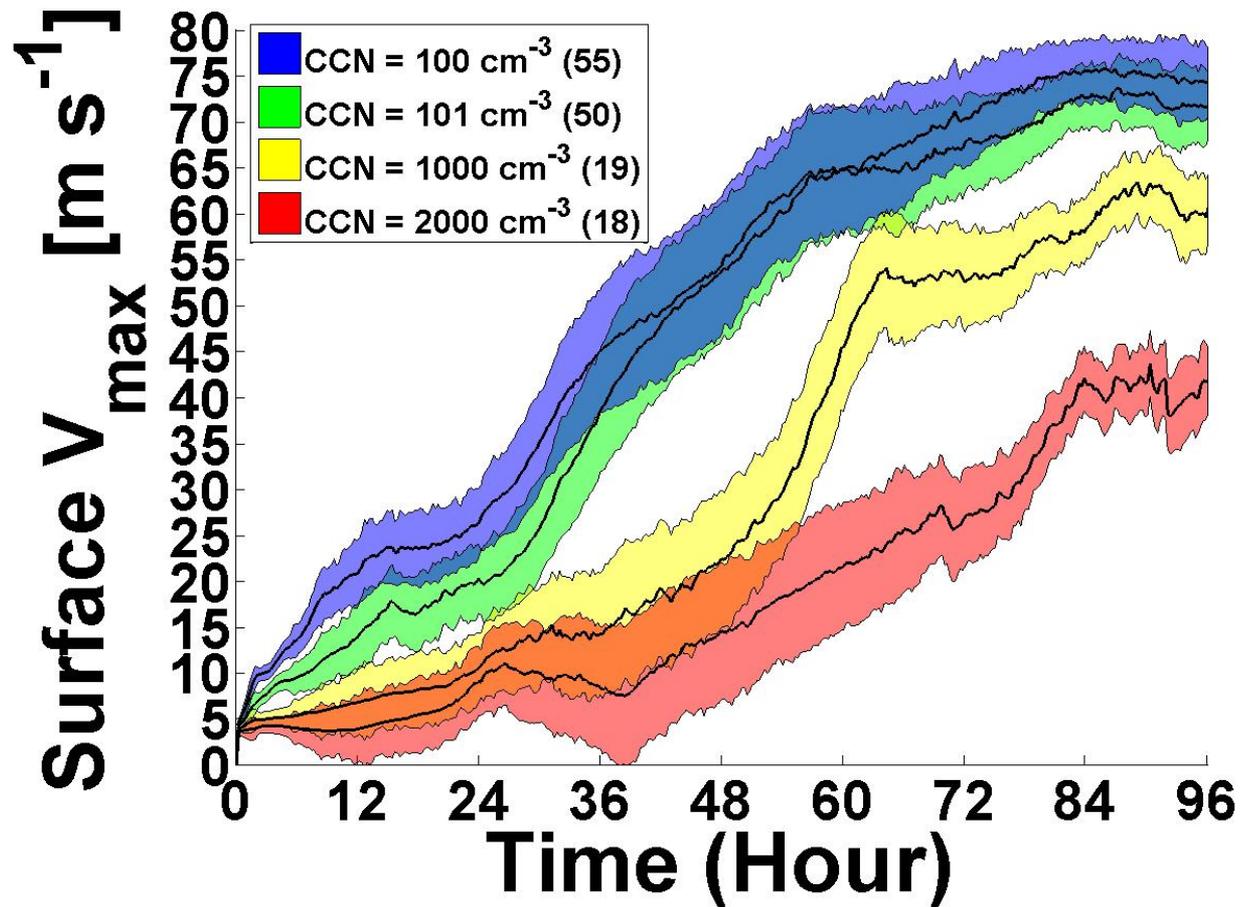


Ensemble where CCN concentration varied for different initial T, q, RMW, vortex height, bubble temperature



Ensemble where CCN concentration varied for different initial T, q, RMW, vortex height, bubble temperature

Distinct monotonic response to varying CCN



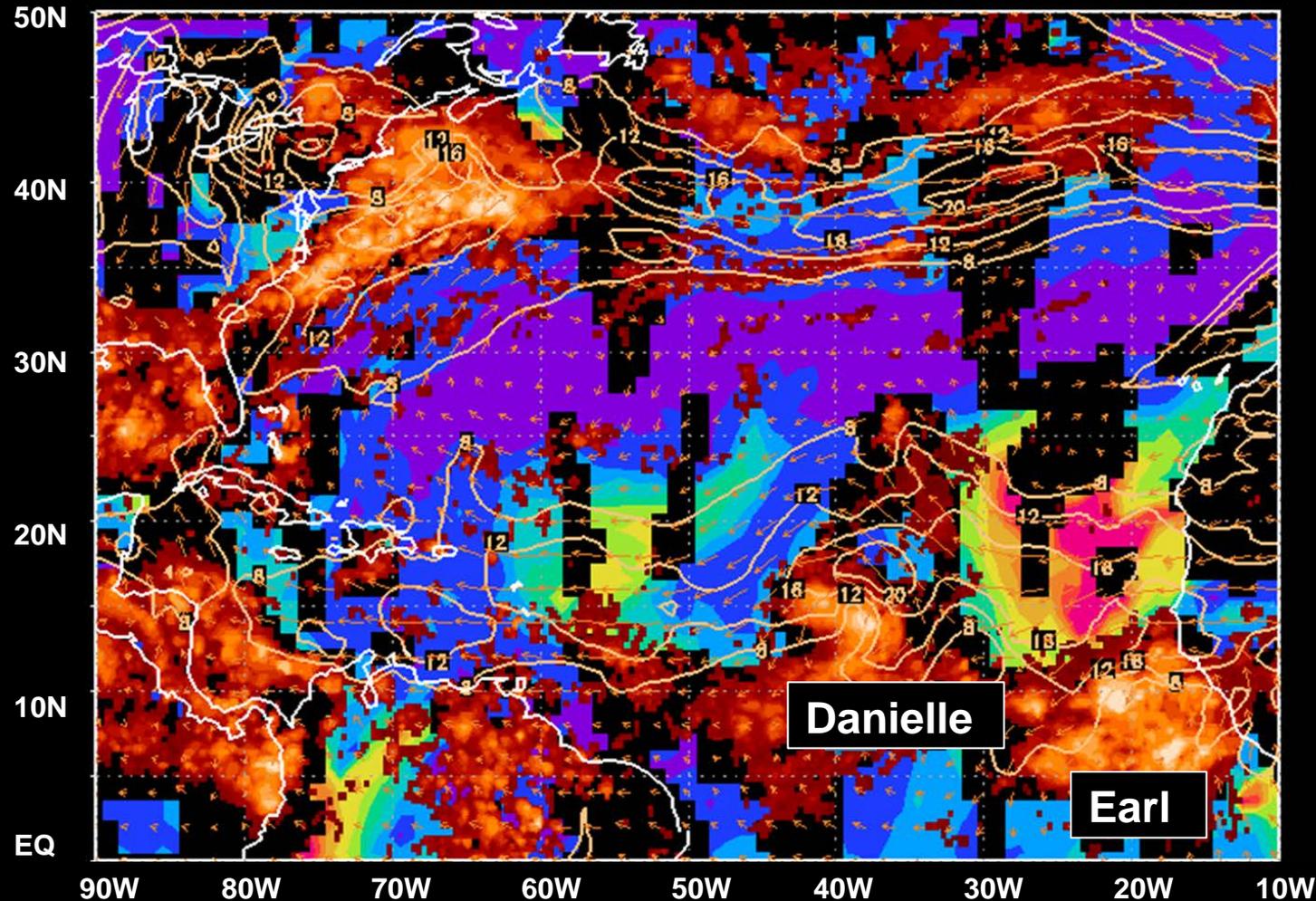
Ensemble where CCN concentration varied for different initial T, q, RMW, vortex height, bubble temperature

Distinct monotonic response to varying CCN

Future: dependence on time of CCN input, response to IN

Evolution of Danielle, Earl, and Fiona (2010)

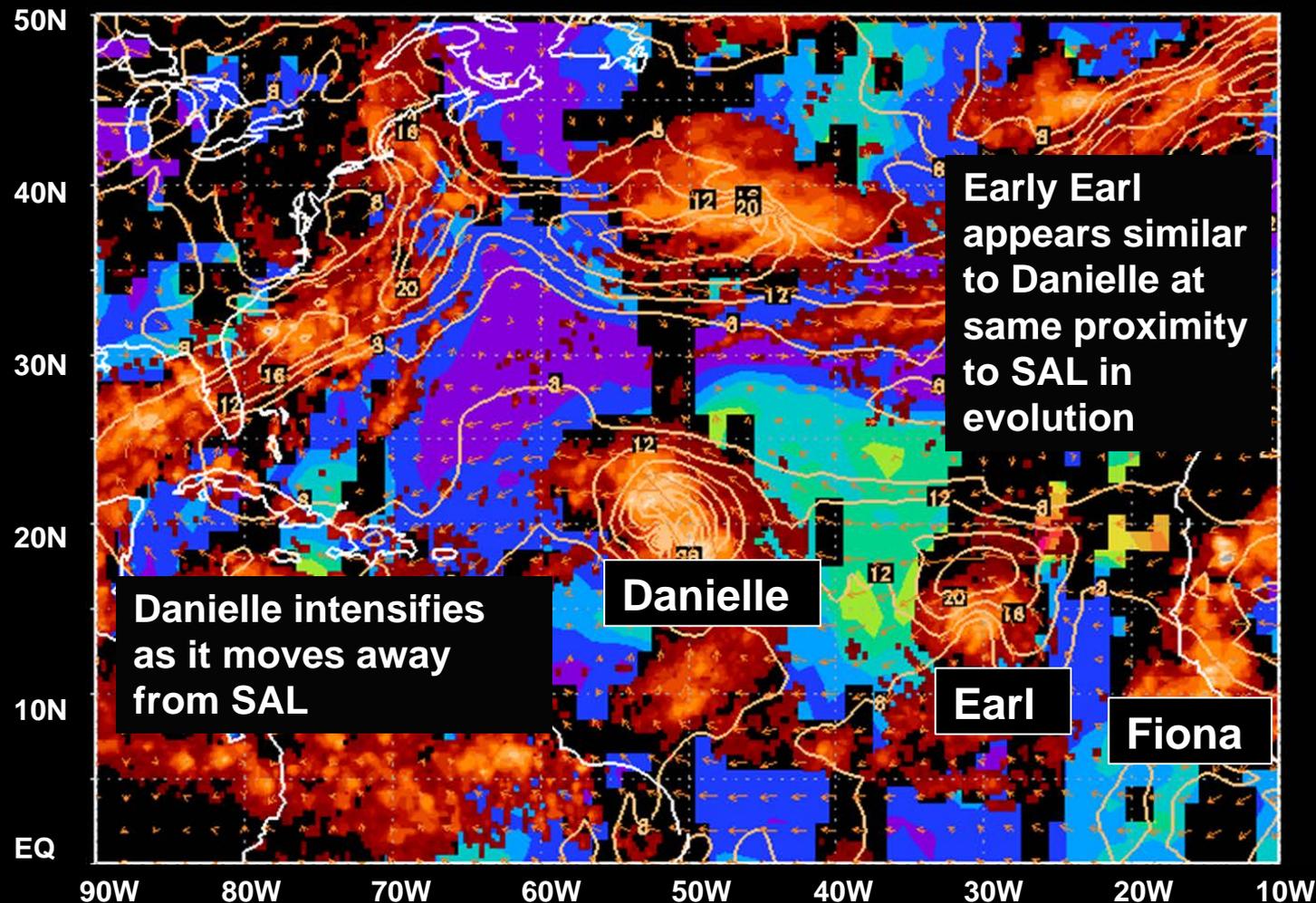
23 August 2010 (700 hPa wind, AOD, 24-hr rainfall)



- Orange coloring denotes 24 hour TRMM rainfall
- Color-filled contours represent AOD
- Contoured is 700 hPa wind

Evolution of Danielle, Earl, and Fiona (2010)

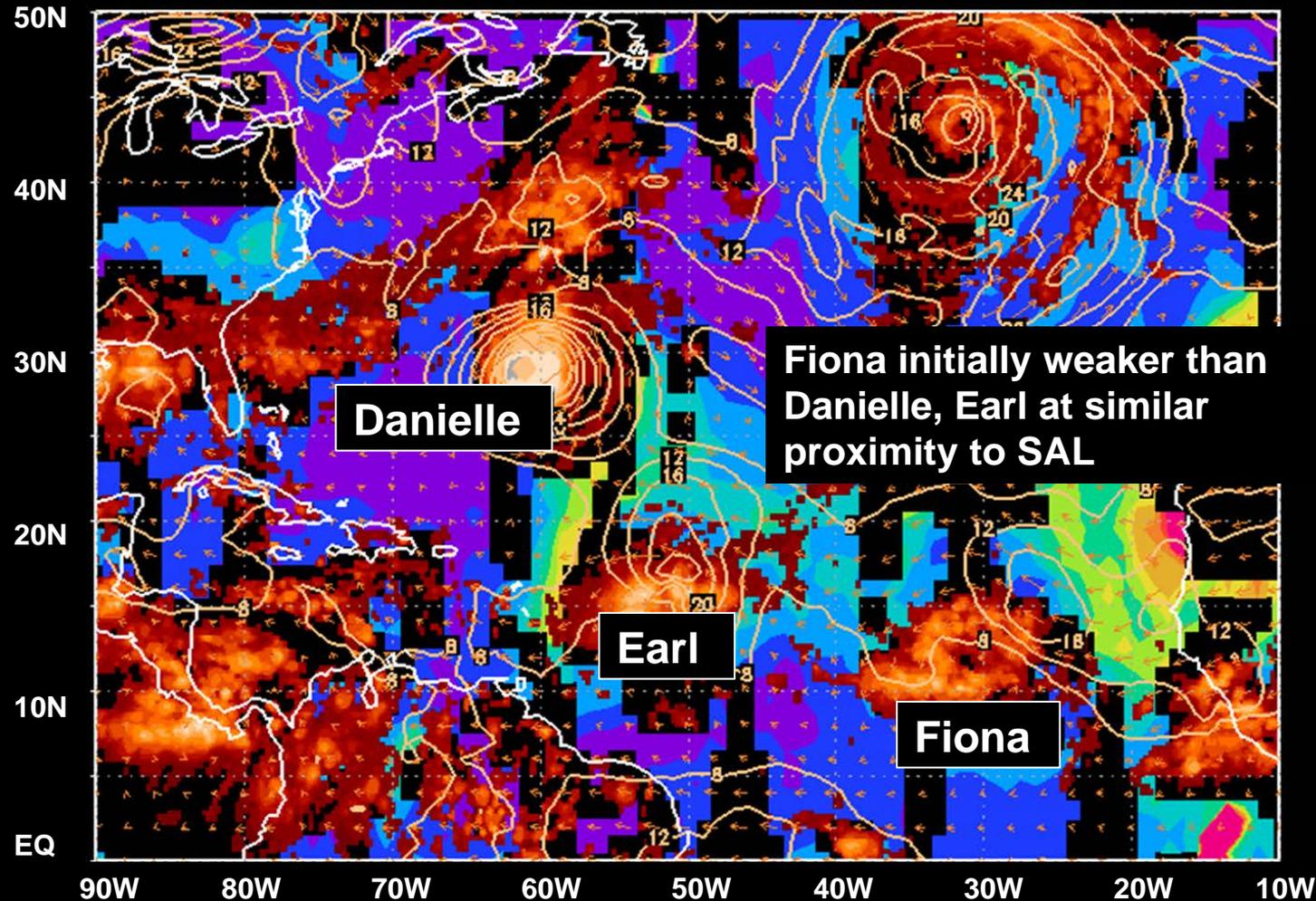
25 August 2010 (700 hPa wind, AOD, 24-hr rainfall)



- Orange coloring denotes 24 hour TRMM rainfall
- Color-filled contours represent AOD
- Contoured is 700 hPa wind

Evolution of Danielle, Earl, and Fiona (2010)

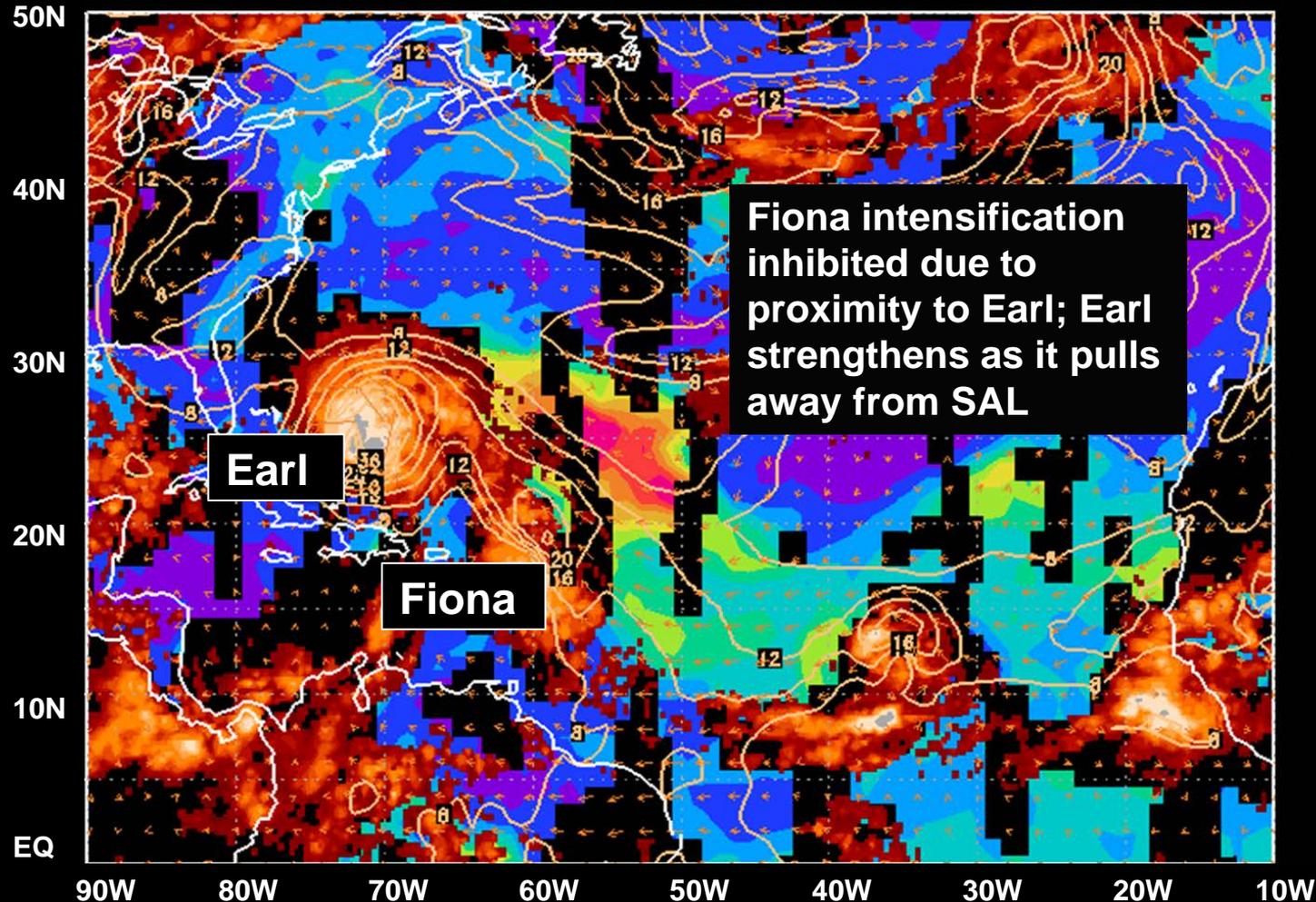
28 August 2010 (700 hPa wind, AOD, 24-hr rainfall)



- Orange coloring denotes 24 hour TRMM rainfall
- Color-filled contours represent AOD
- Contoured is 700 hPa wind

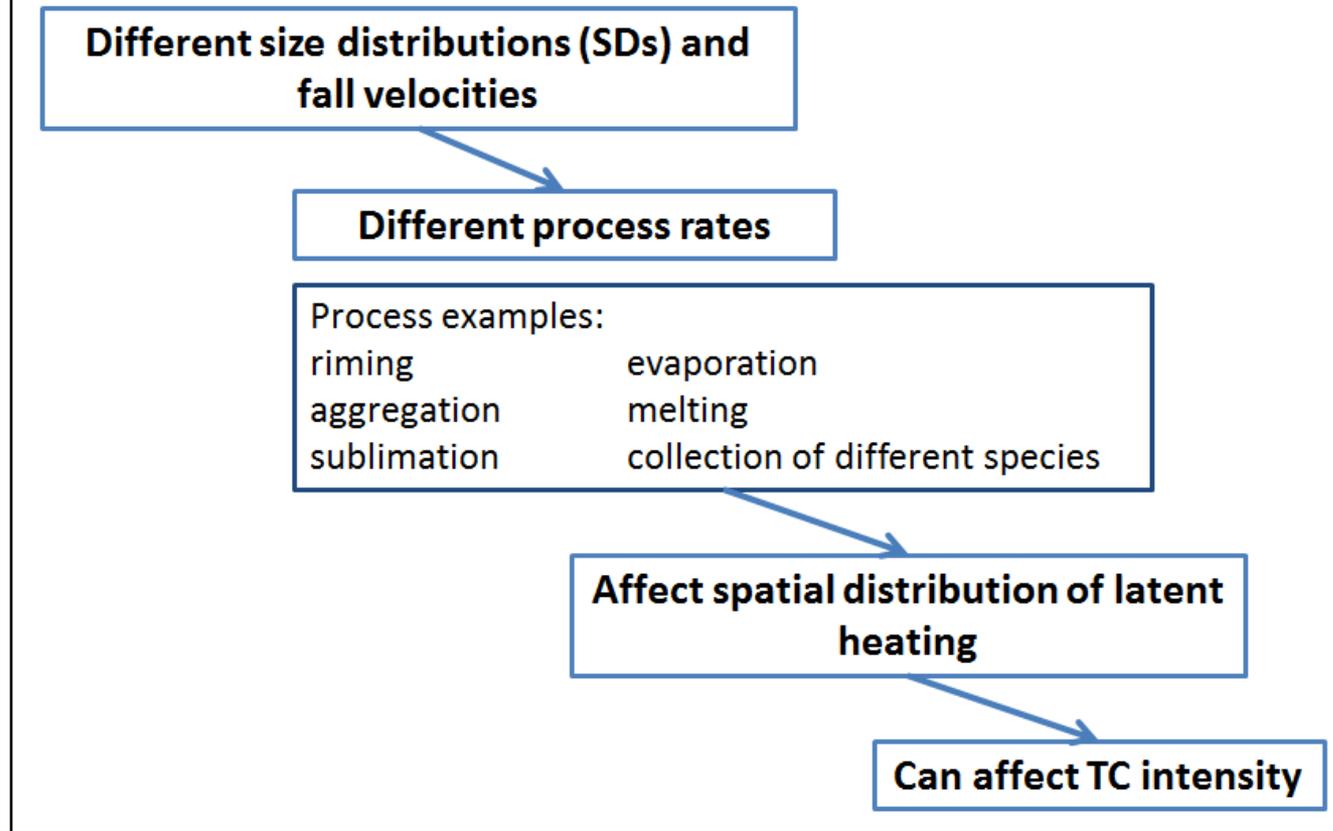
Evolution of Danielle, Earl, and Fiona (2010)

01 September 2010 (700 hPa wind, AOD, 24-hr rainfall)

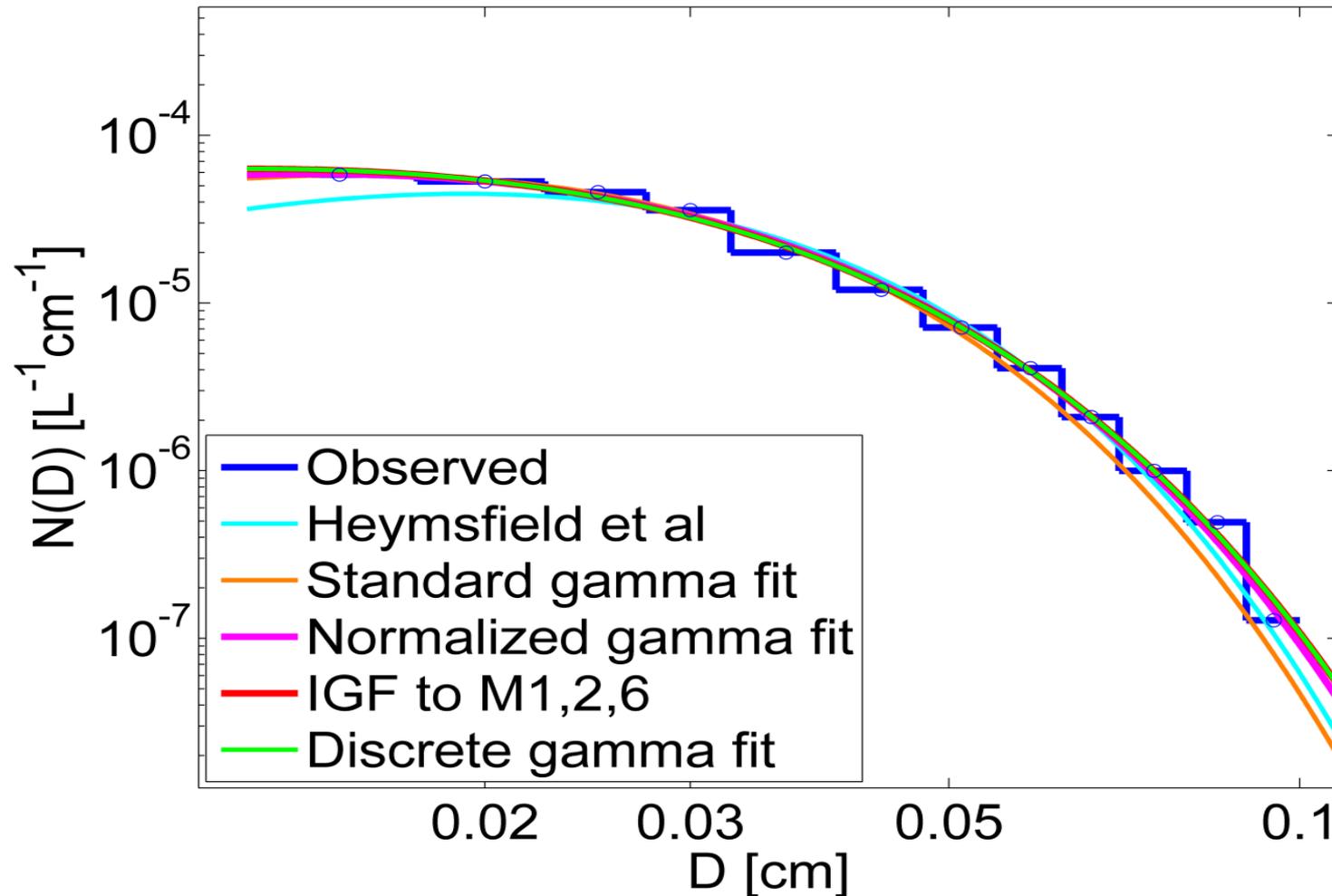


- Orange coloring denotes 24 hour TRMM rainfall
- Color-filled contours represent AOD
- Contoured is 700 hPa wind

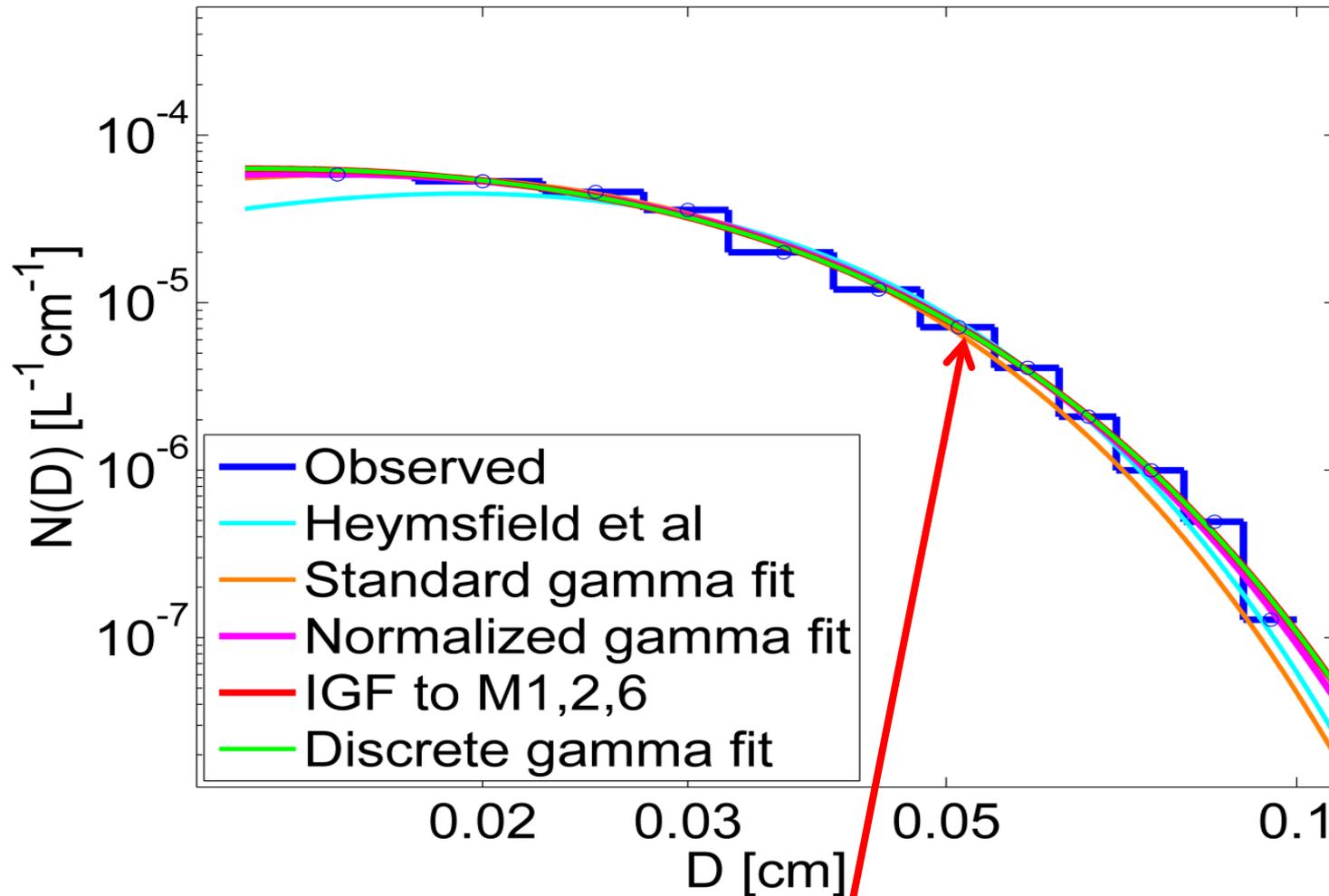
Cloud Microphysics Affects TC Intensity



- Gamma functions $N(D) = N_0 D^\mu e^{-\lambda D}$ are used to represent SDs
How do N_0 , μ and λ vary with environmental/cloud conditions, and how accurately can they be determined?

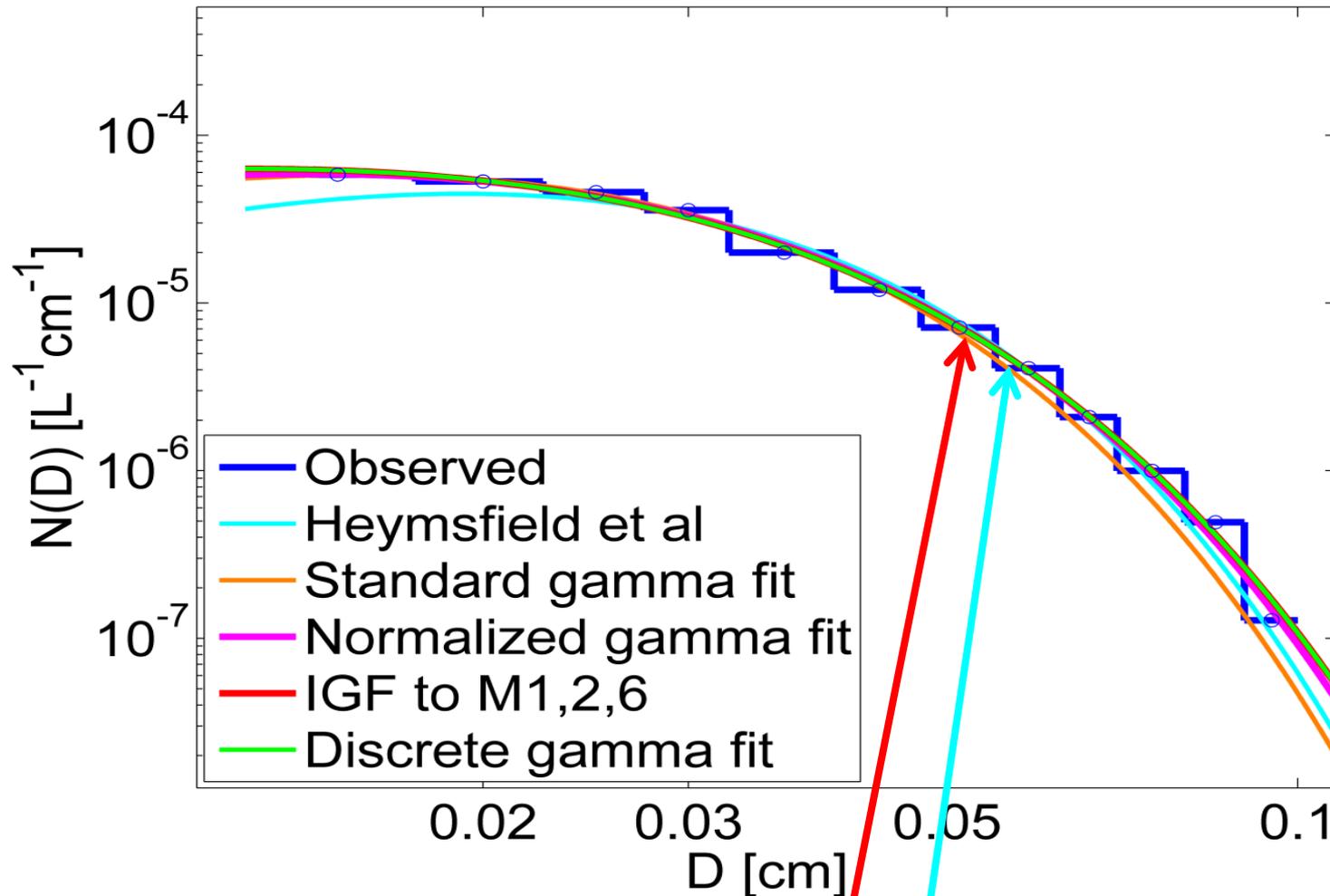


Even though fits all look quite good, there can be huge range in N_0 , λ and μ



Even though fits all look quite good, there can be huge range in N_0 , λ and μ

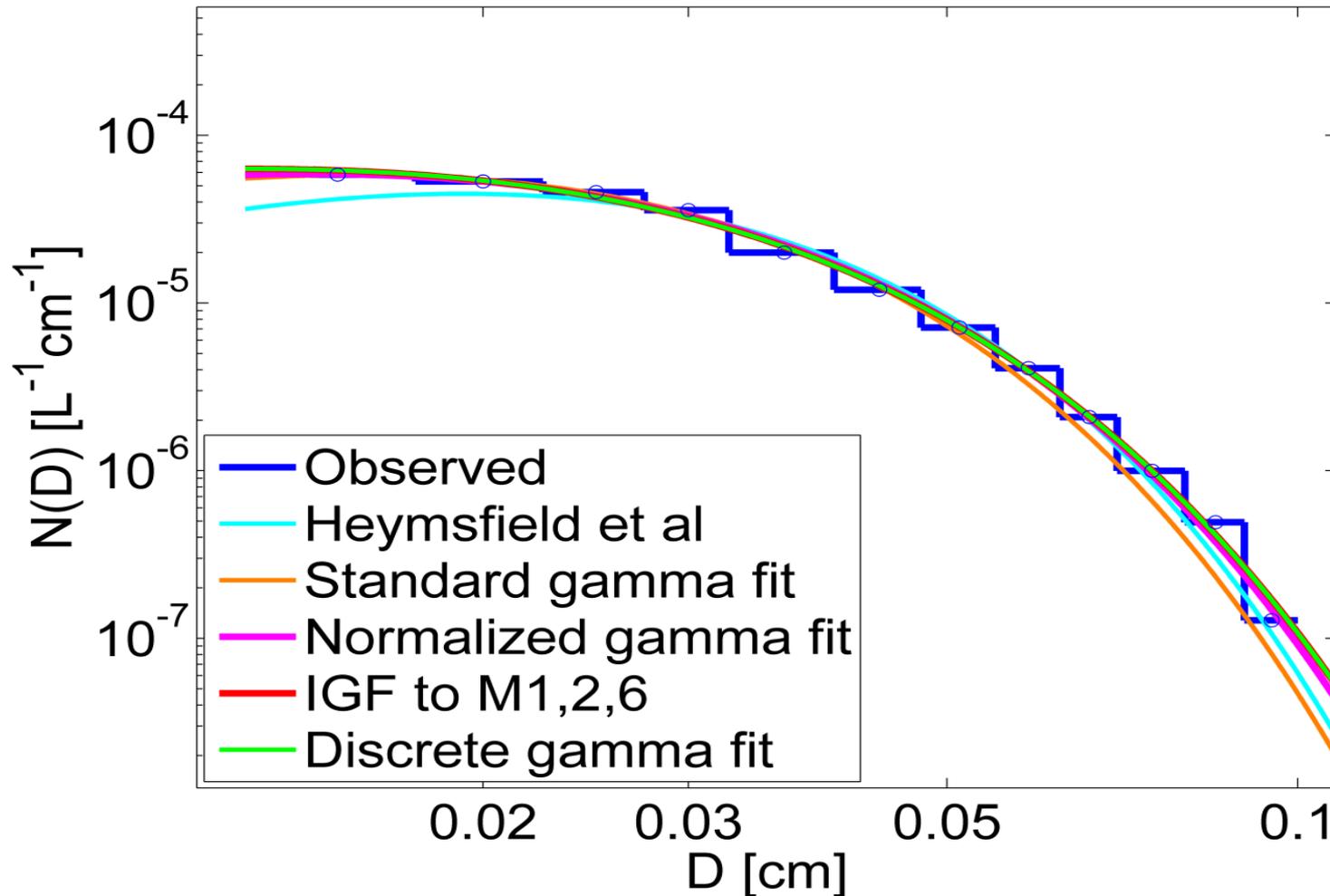
IGF: $N_0 = 6.3 \times 10^{-1}$ $\mu = 1.86$; $\lambda = 1.1 \times 10^2$



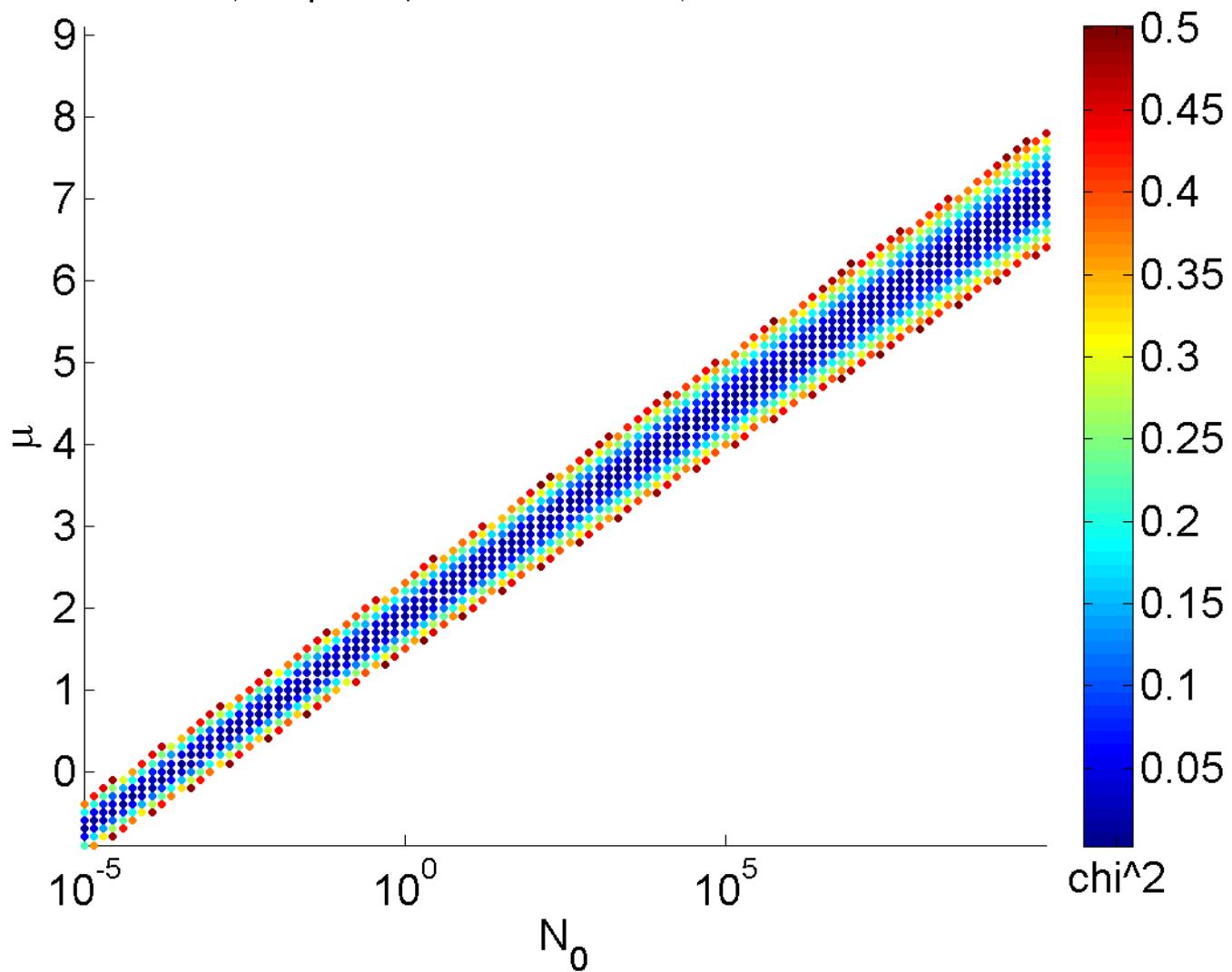
Even though fits all look quite good, there can be huge range in N_0 , λ and μ

IGF: N_0 6.3×10^{-1} $\mu=1.86$; $\lambda = 1.1 \times 10^2$

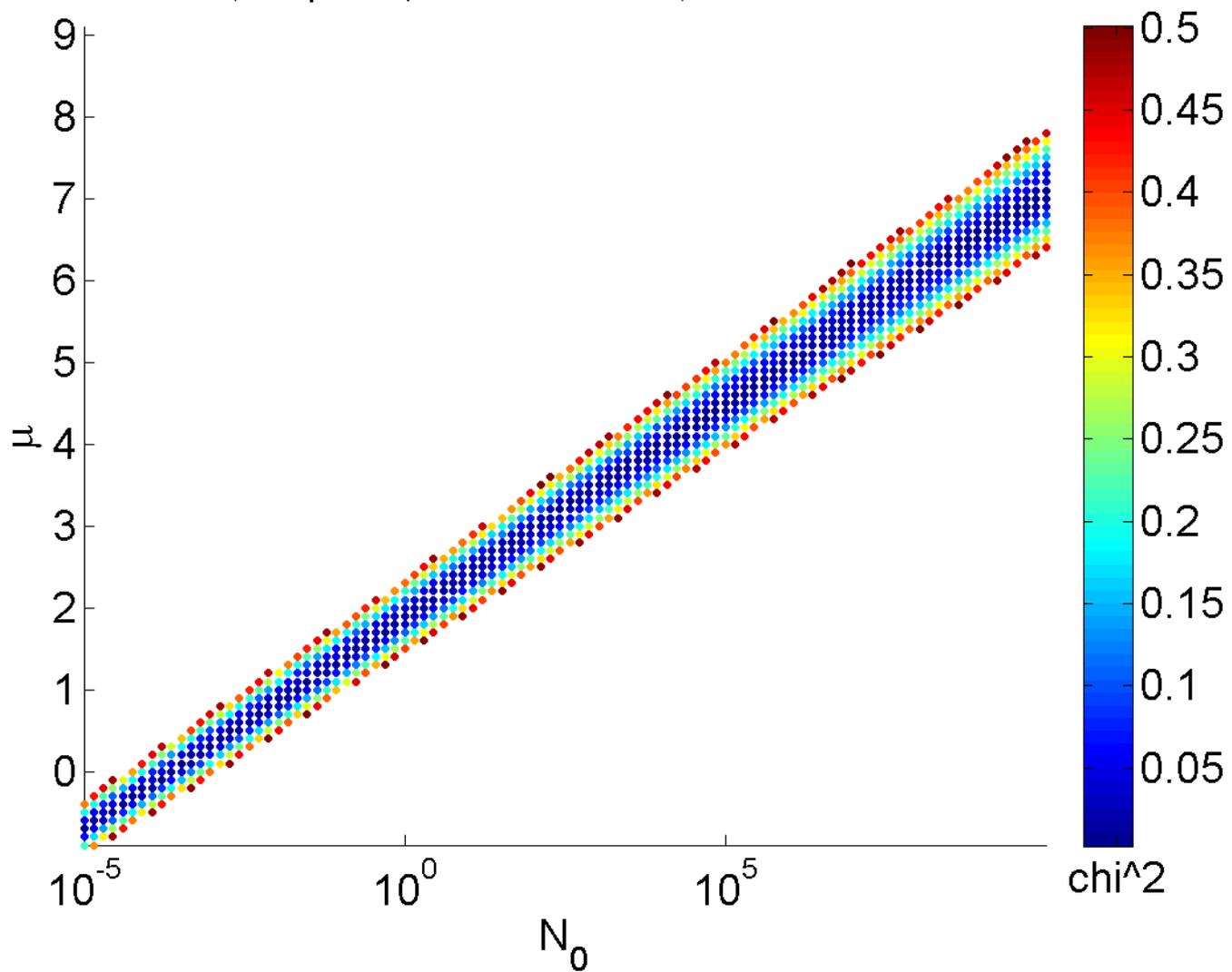
Hey: N_0 7.3×10^1 $\mu=2.97$; $\lambda = 1.4 \times 10^2$



There is broad range of $N_0/\mu/\lambda$ that fit SD well
 $N_0/\mu/\lambda$ determined depend on tolerance
allowed
→ Can't represent by single $N_0/\mu/\lambda$ value



**Surface of equally realizable solutions
with $\Delta\chi^2 = 0.5$ of χ_{\min}^2**

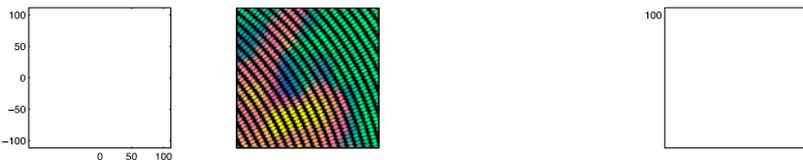
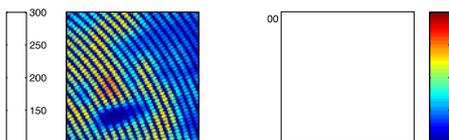


Show how surfaces vary with T, w, TC stage of development in Mascio et al. (2012)

How can operational passive microwave imagery help in observing, understanding and predicting RI processes and pathways?

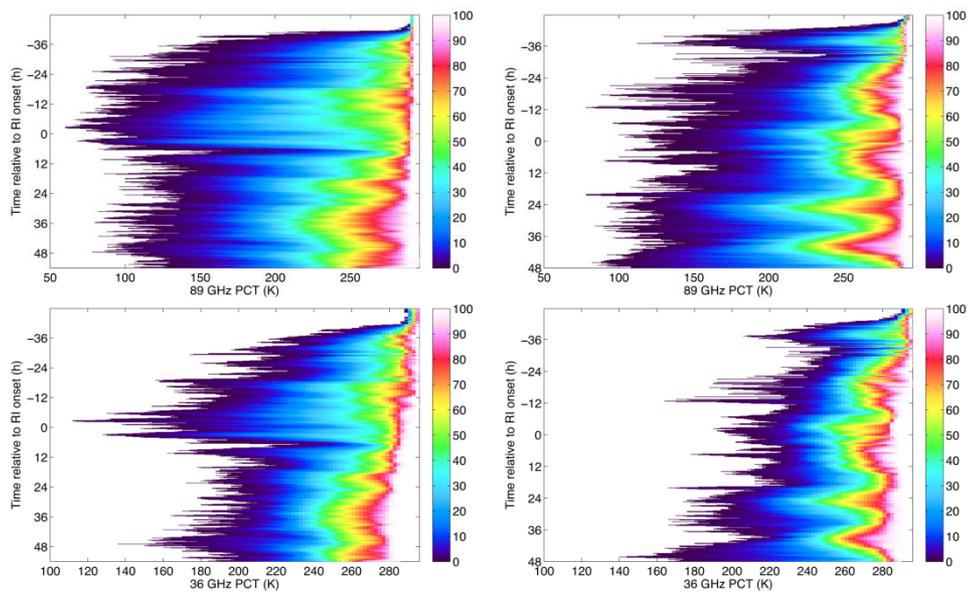
Tool: WRF + Nagoya/Goddard Satellite Data Simulator Unit

Earl 6 h prior to RI Ike



Use simulations as prototypes for development of RI prediction algorithms, understand processes during RI going on in the model, and compare against FC, satellite observations

Earl T_b CDFs Ike

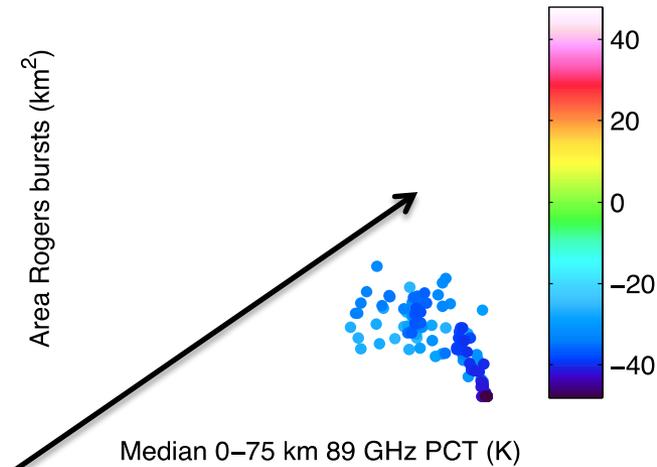
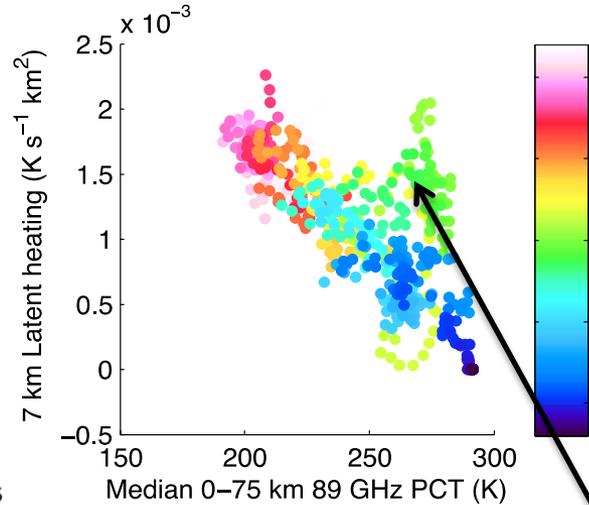
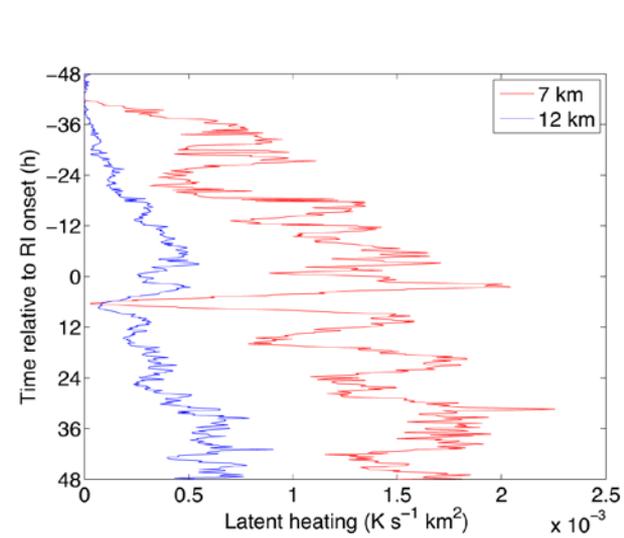


Earl – Gradual deepening, Ike – Rapid deepening +18-24 h

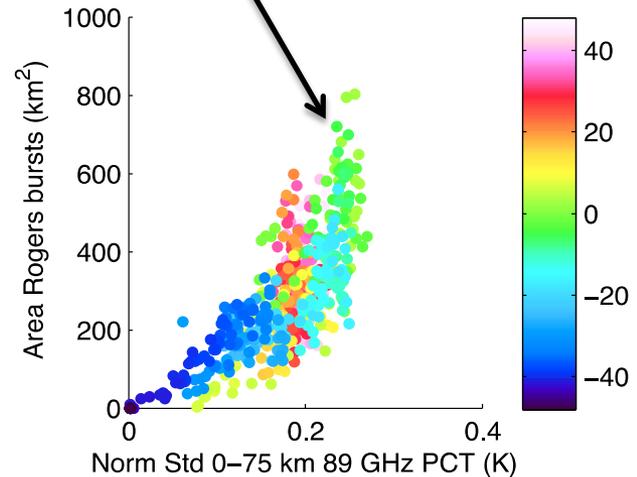
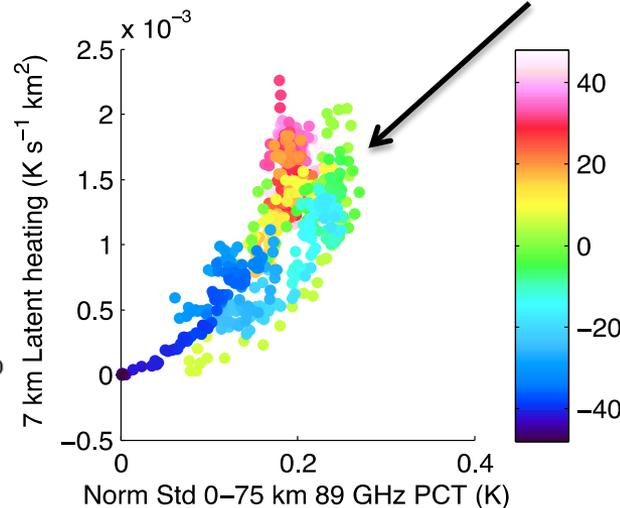
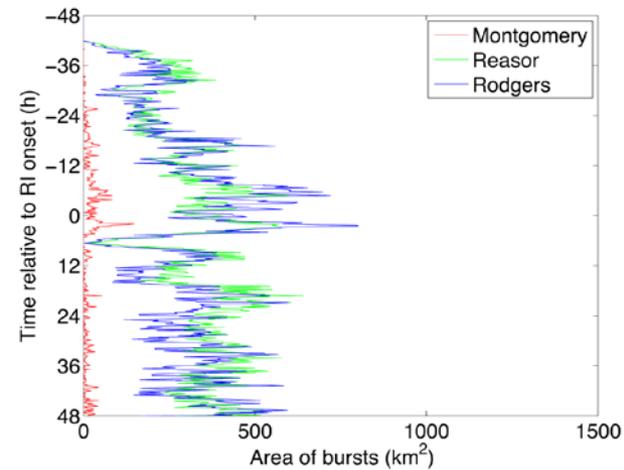
Earl WRF+SDSU output

Simulated latent heating and convective burst activity relative to RI onset

89 GHz Tb statistics (0-75 km median [top] and std dev [bottom]) as a function of time relative to RI onset (colors)

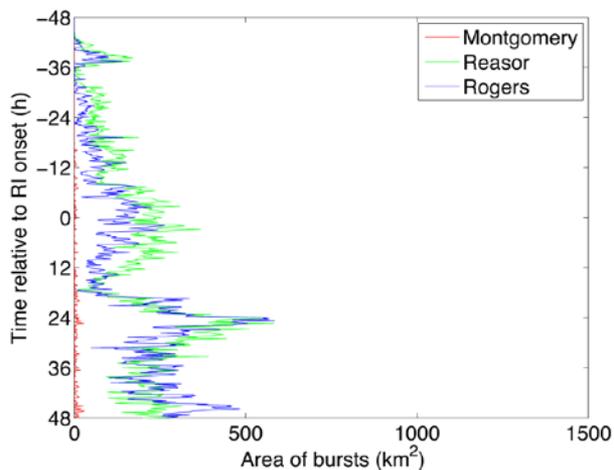
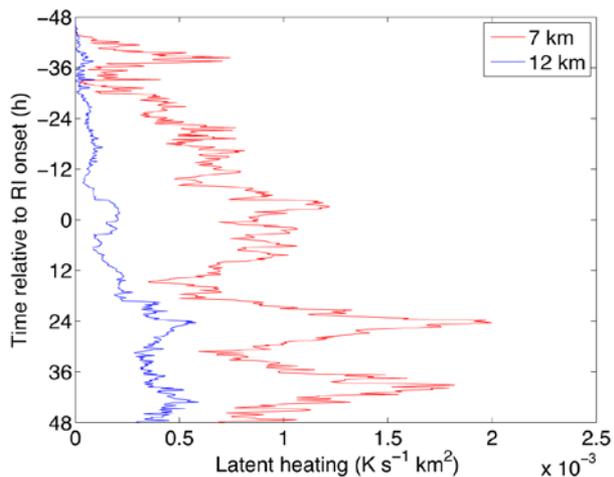


Asymmetric, "bursty" convection (modest median, large std dev) 18-0 h prior to RI

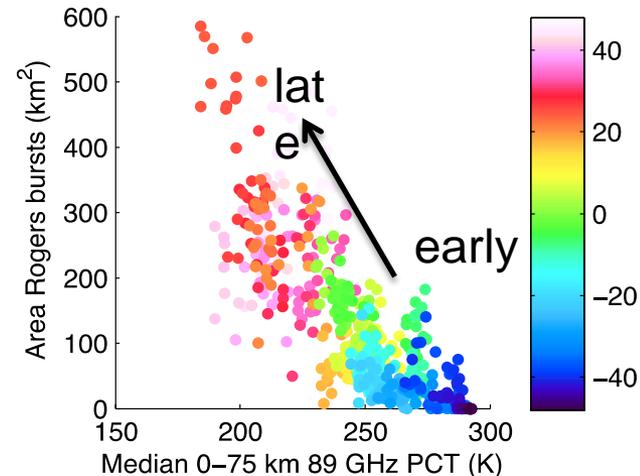
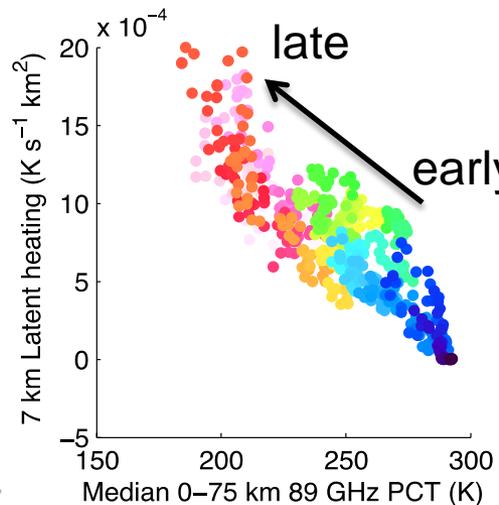


Like WRF+SDSU output

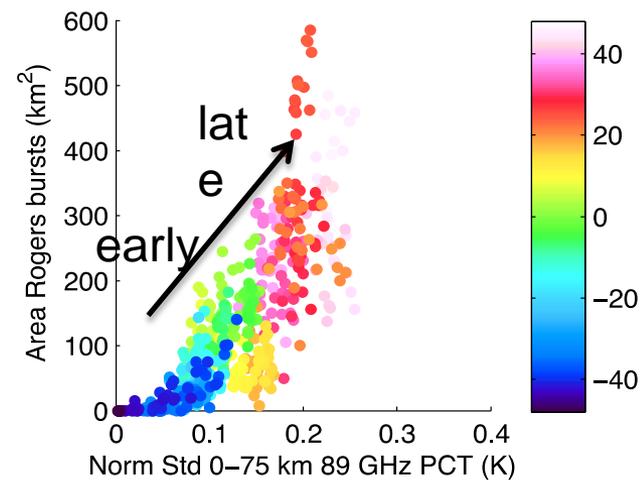
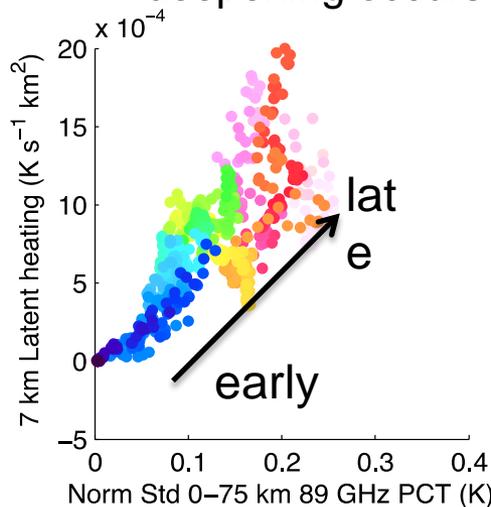
Simulated latent heating and convective burst activity relative to RI onset



89 GHz Tb statistics (0-75 km median [top] and std dev [bottom]) as a function of time relative to RI onset (colors)



RI phase begins with symmetric modest convection, rapid deepening occurs 18-30 h after RI onset bursts ensue



See more at Poster Session II



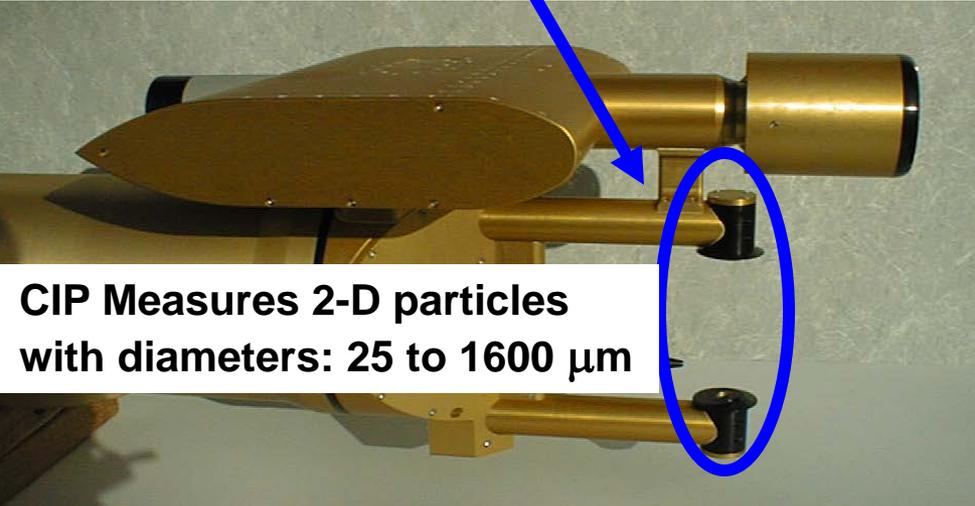
Summary

- **WRF simulations of Dennis, Earl and Ike showing mid-level increases in vertical velocity outliers and hydrometeor upward fluxes are precursors to RI**
 - Next step is to compare against GRIP+HS3 observations
- **Techniques for understanding uncertainties in gamma parameters derived from in-situ data available**
 - Next step is to apply to GRIP in-situ data
- **Increases in CCN appear to weaken TC from idealized simulation**
 - Next case is to simulate Earl and evaluate against observations
- **Necessary to examine causes of RI as function of shear**
 - Look for consistencies between model & 2010/2012-14 observations

Probes

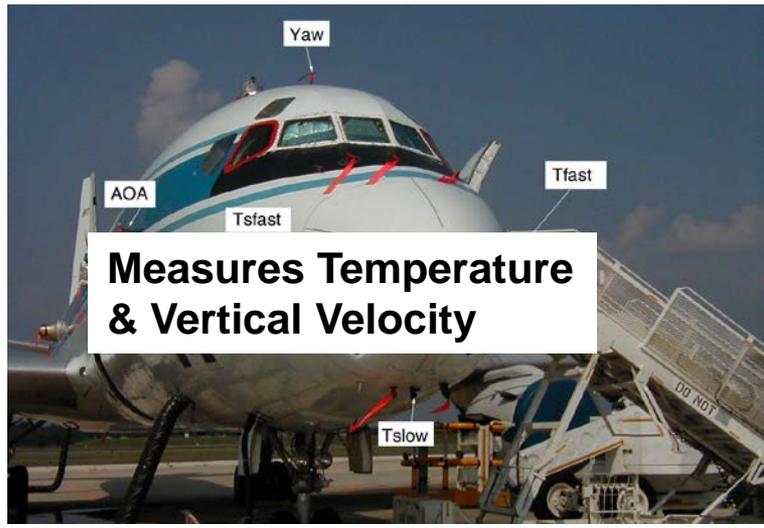
Cloud, Aerosol, and Precipitation Spectrometer (CAPS)

- *Cloud and Aerosol Spectrometer (CAS)*
- *Cloud Imaging Probe (CIP)*



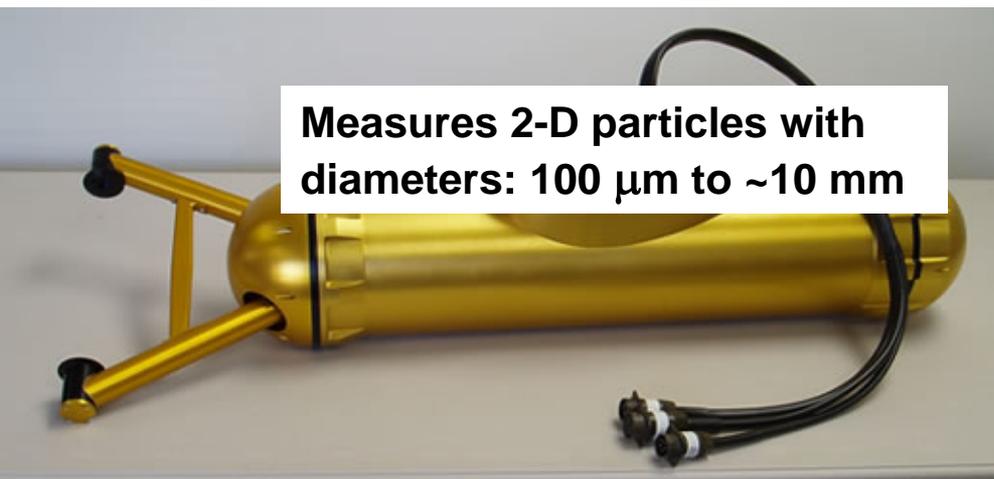
CIP Measures 2-D particles with diameters: 25 to 1600 μm

Meteorological Measurement System (MMS)



Measures Temperature & Vertical Velocity

Precipitation Imaging Probe (PIP)

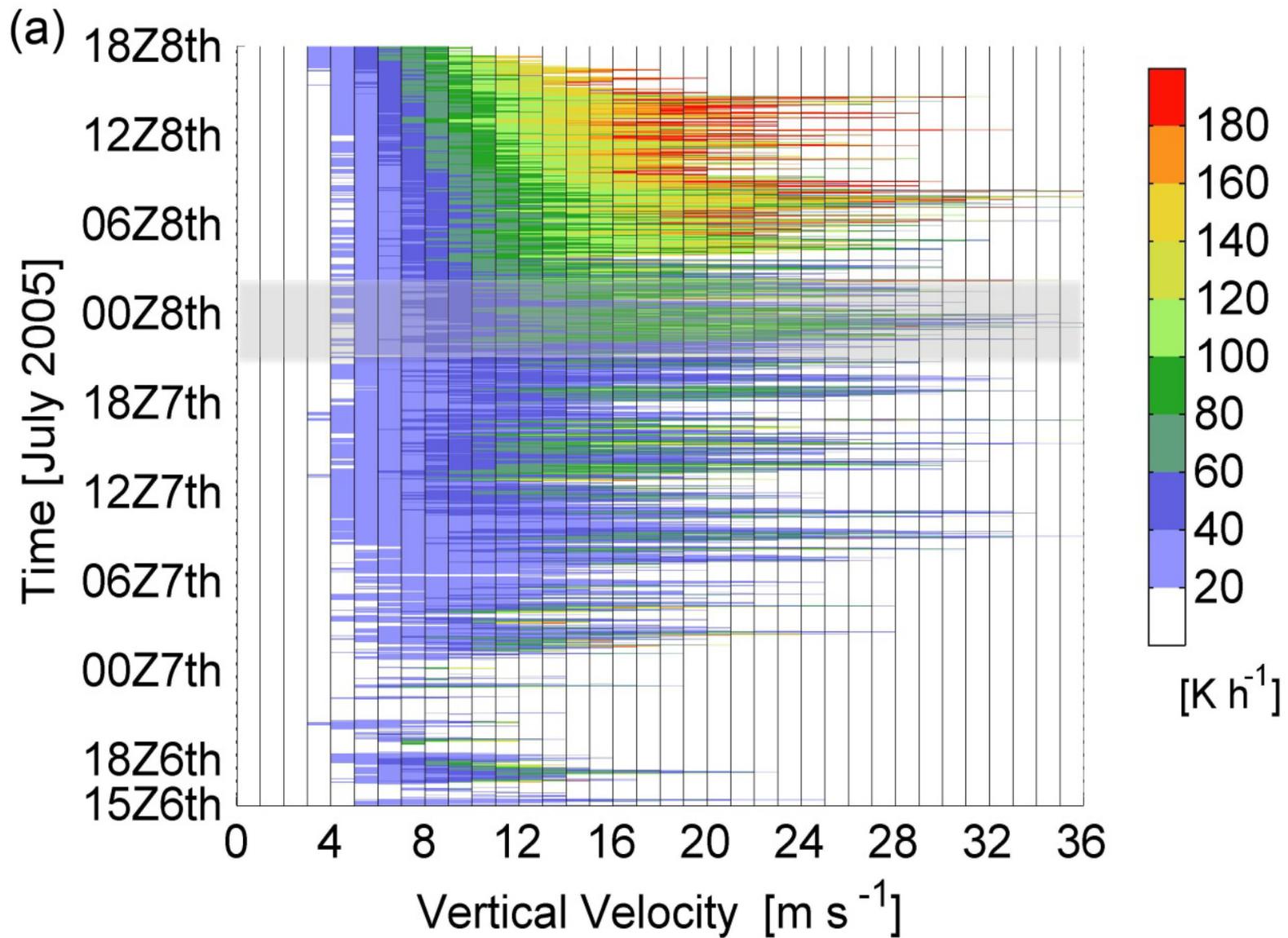


Measures 2-D particles with diameters: 100 μm to ~10 mm

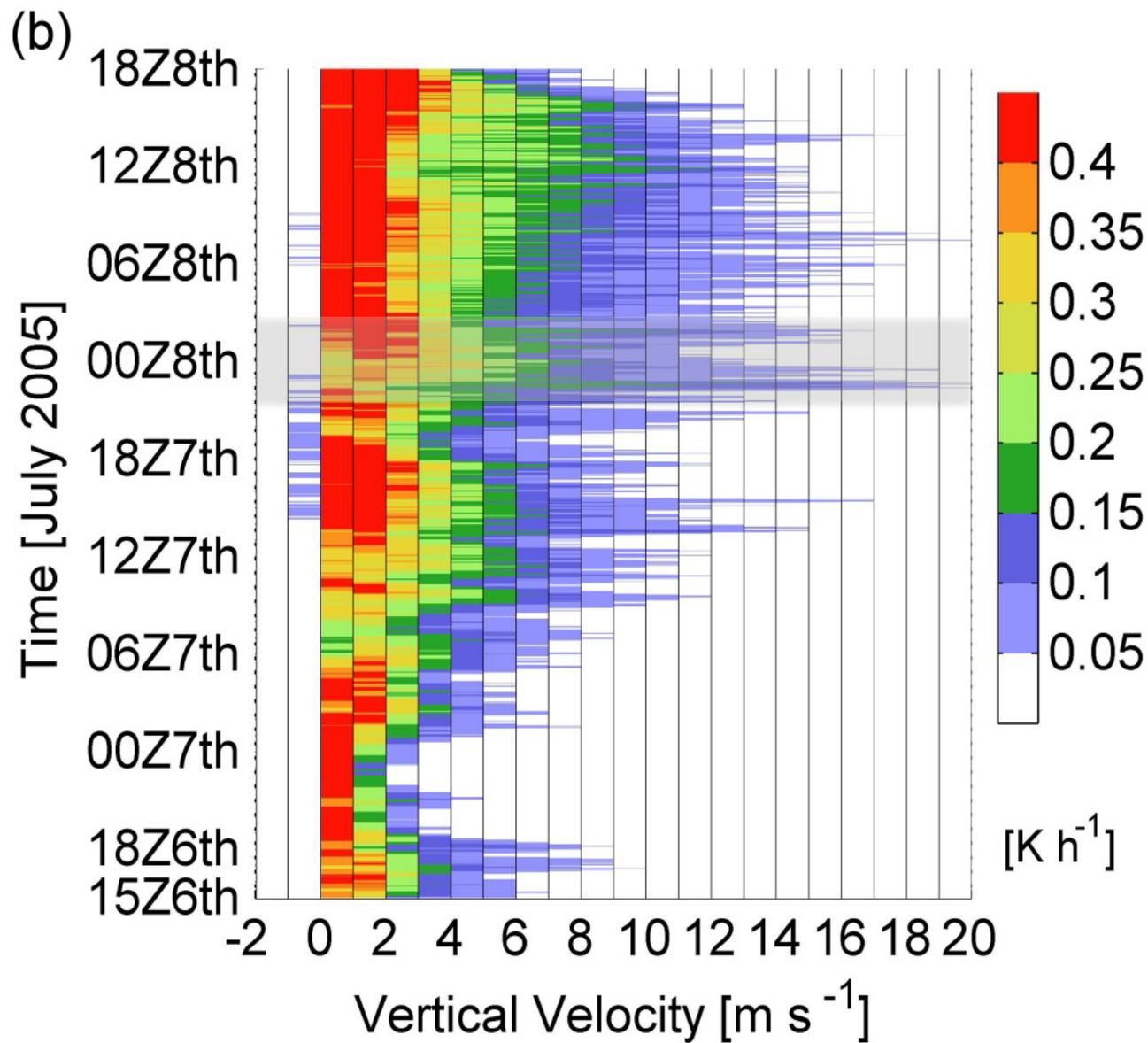
Composite Size Distribution:

CIP: $125 < D < 1000 \mu\text{m}$

PIP: $D > 1000 \mu\text{m}$



Average latent heating in volume elements with large vertical velocity is large, and increases after RI



BUT, lower velocity bins dominate aggregate latent heating both before and after RI

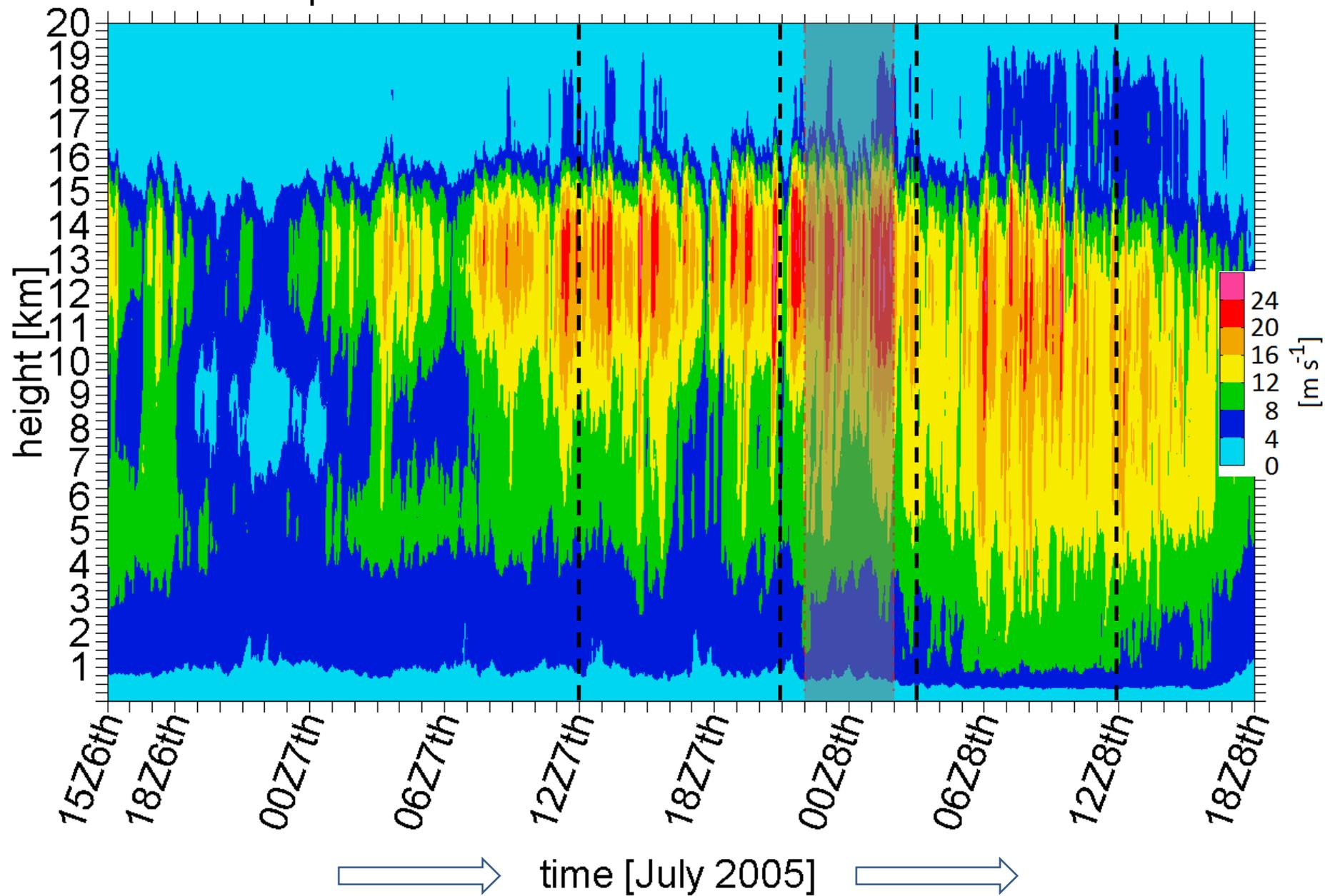


Conclusions from WRF simulations

- **Outliers of w better indicator of RI than averages**
 - Precursor to RI *at upper levels* (e.g., 14 km)
 - Continual broadening and convergence toward TC center
- **Outlier (e.g., 99.9th percentile) w *at lower levels* (e.g., 6 km) increase only *after* onset of and *during* RI**
- **Immediately prior to RI, increase in upward hydrometeor mass flux for $10 < z < 15$ km**
 - Levels with $z < 1.5$ km make big contributions to air mass updraft flux, but this does not seem to be precursor to RI
- **Vertical velocities in 90-99.99th percentile have minimal effect on aggregate latent heating**
 - Latent heating dominated by $w < 2 \text{ m s}^{-1}$
- **Necessary to examine causes of RI as function of shear**

99th percentile **VERTICAL VELOCITY**

$r < 75$ km



Average VERTICAL VELOCITY

$r < 75$ km

